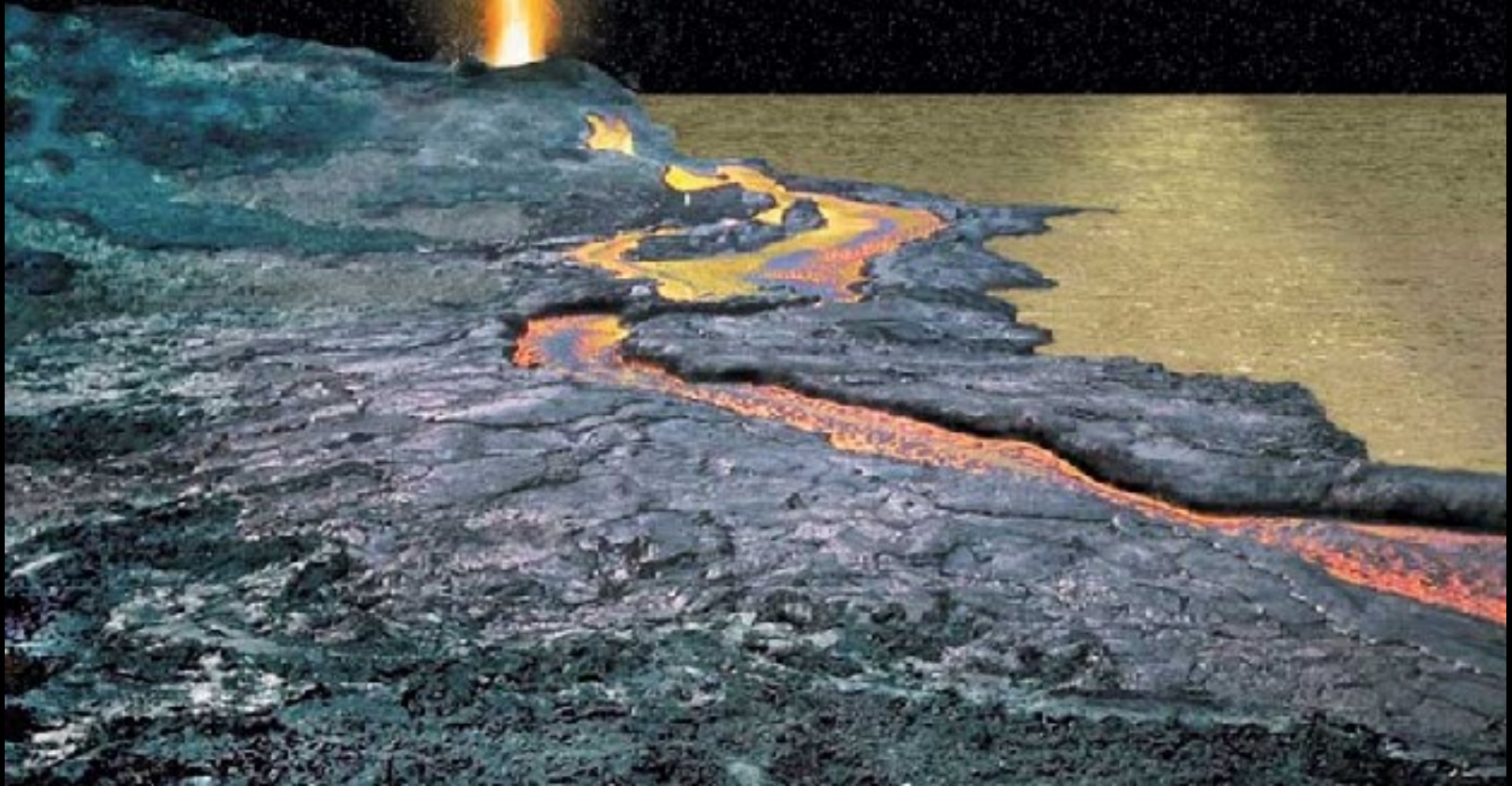
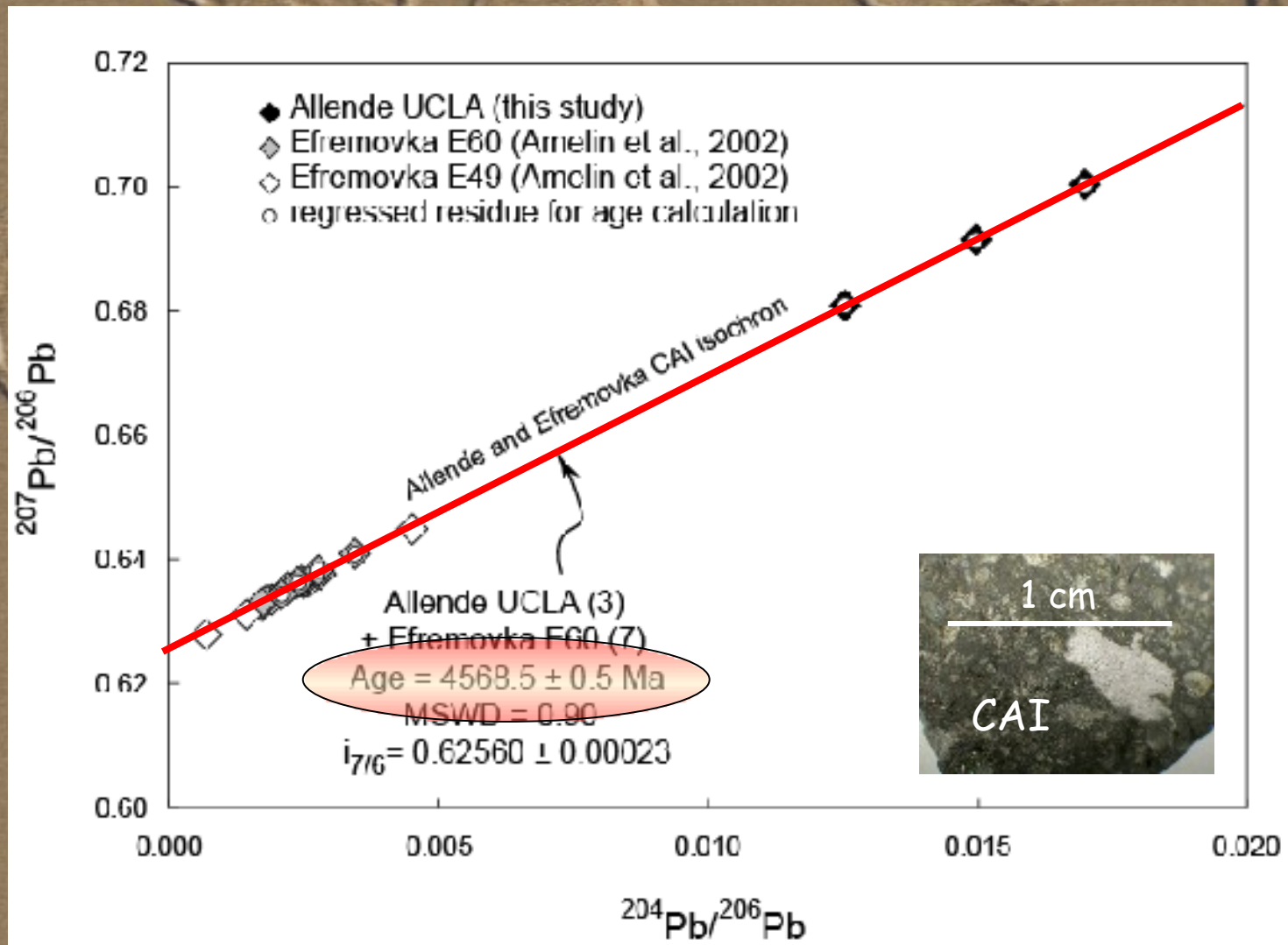


L'âge de la première croûte
continentale et le début de
la tectonique des plaques



Tout a commencé il y à 4568,5 Ma!



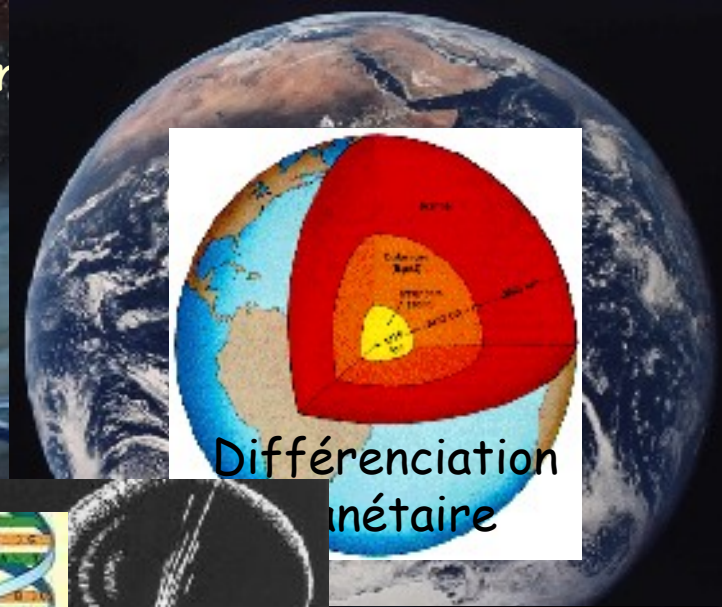
La nébuleuse



L'accrétion



Corps plan



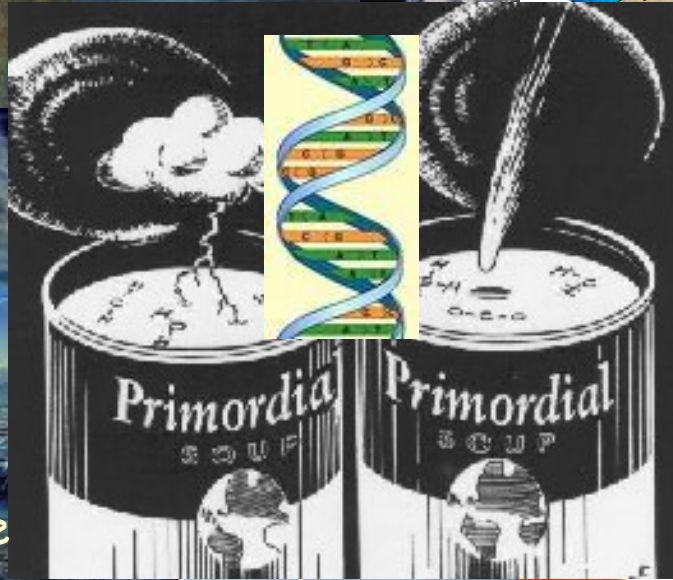
Différenciation planétaire

Dynamos/cha

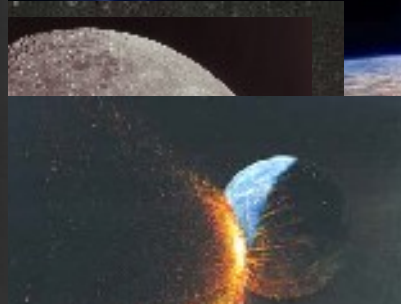


Dégazage/atmosphère

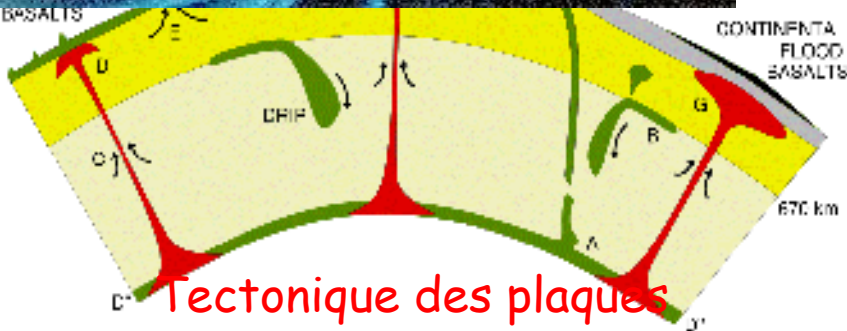
© 2002 Don Dixon - cosmoglyphics.com



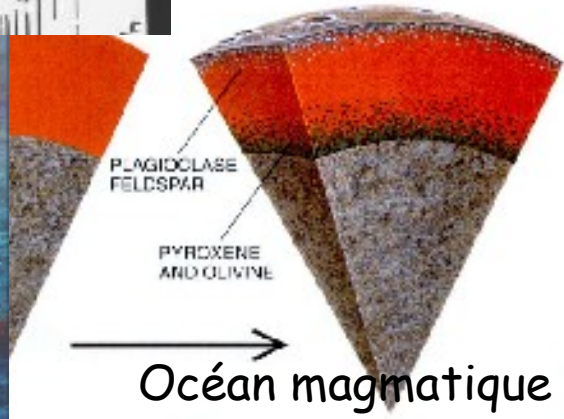
Bombardement



continents



Tectonique des plaques



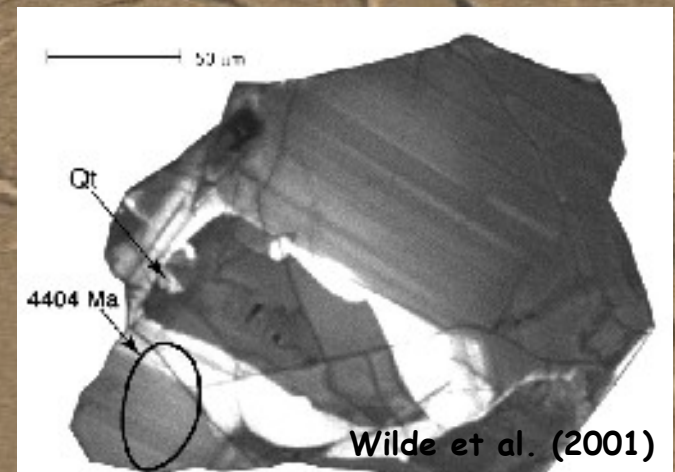
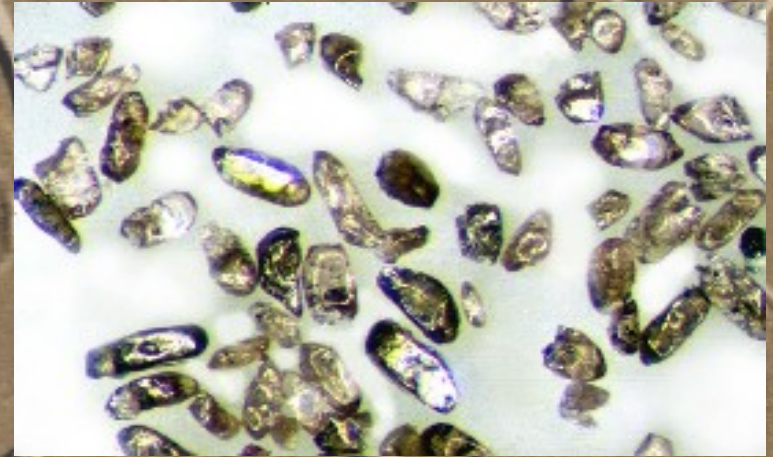
Océan magmatique

Un des grandes questions sur
l'évolution de la Terre:
A quel moment sont apparus les
premiers continents?

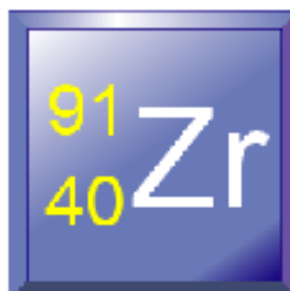
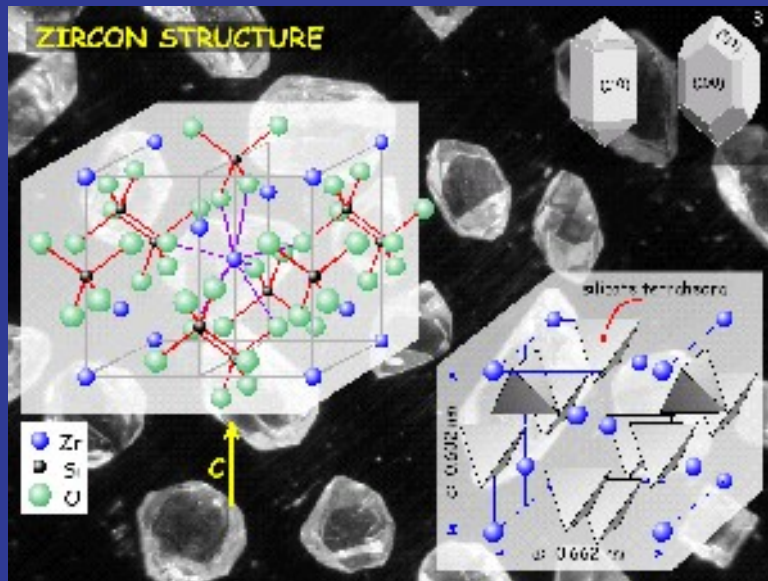


L'outil idéal: les zircons....

...petits, mais résistants, à l'heure et gèle l'Hf!



ZrSiO₄



abundance: ~10 ppm
(silicate Earth)

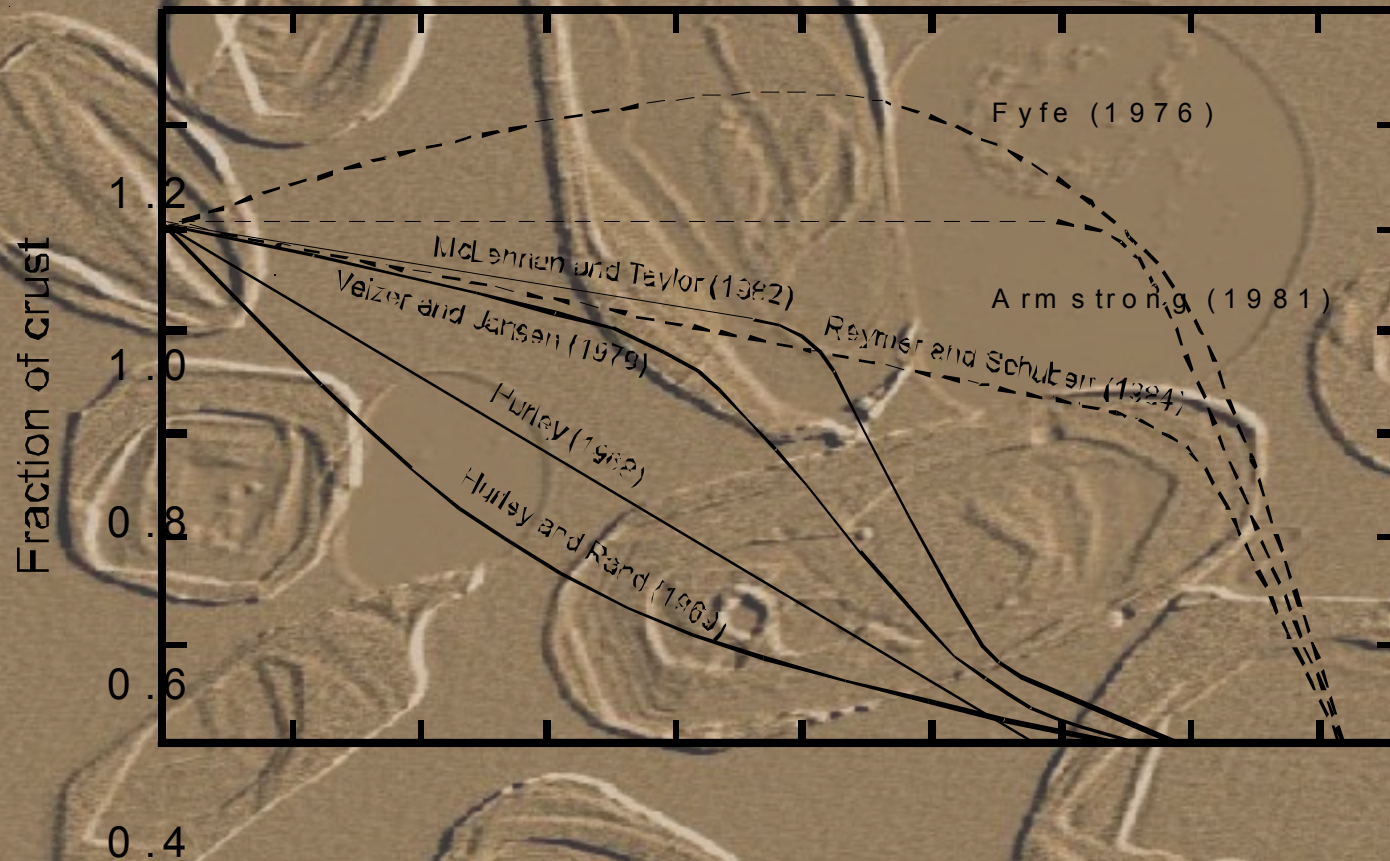
valence state: +4

ionic radius: 0.92 Å

	IA	IIA		IVB	VIB								IVA	VIA	VIIIA				
				IVB	VIB								III A	VA	VII A				
2	Li 3	Be 4											B 5	C 6	N 7	O 8	F 9	Ne 10	
3	Na 11	Mg 12	IIIB										Al 13	Si 14	P 15	S 16	Cl 17	Ar 18	
4	K 19	Ca 20											Ga 31	Ge 32	As 33	Se 34	Br 35	Kr 36	
5	Rb 37	Sr 38	Y 39	Zr 40	Nb 41	Mo 42	Tc 43	Ru 44	Rh 45	Pd 46	Ag 47	Cd 48	In 49	Sn 50	Sb 51	Te 52	I 53	Xe 54	
6	Cs 55	Ba 56	La 57	Hf 72	Ta 73	W 74	Re 75	Os 76	Ir 77	Pt 78	Au 79	Hg 80	Tl 81	Pb 82	Bi 83	Po 84	At 85	Rn 86	
7	87	88	89																
				Ca 58	Pr 59	Nd 60	Pm 61	Sm 62	Eu 63	Gd 64	Tb 65	Dy 66	Ho 67	Er 68	Tm 69	Yb 70	Lu 71		
				Th 90		U 92													
				91		93	94	95	96	97	98	99	100	101	102	103			

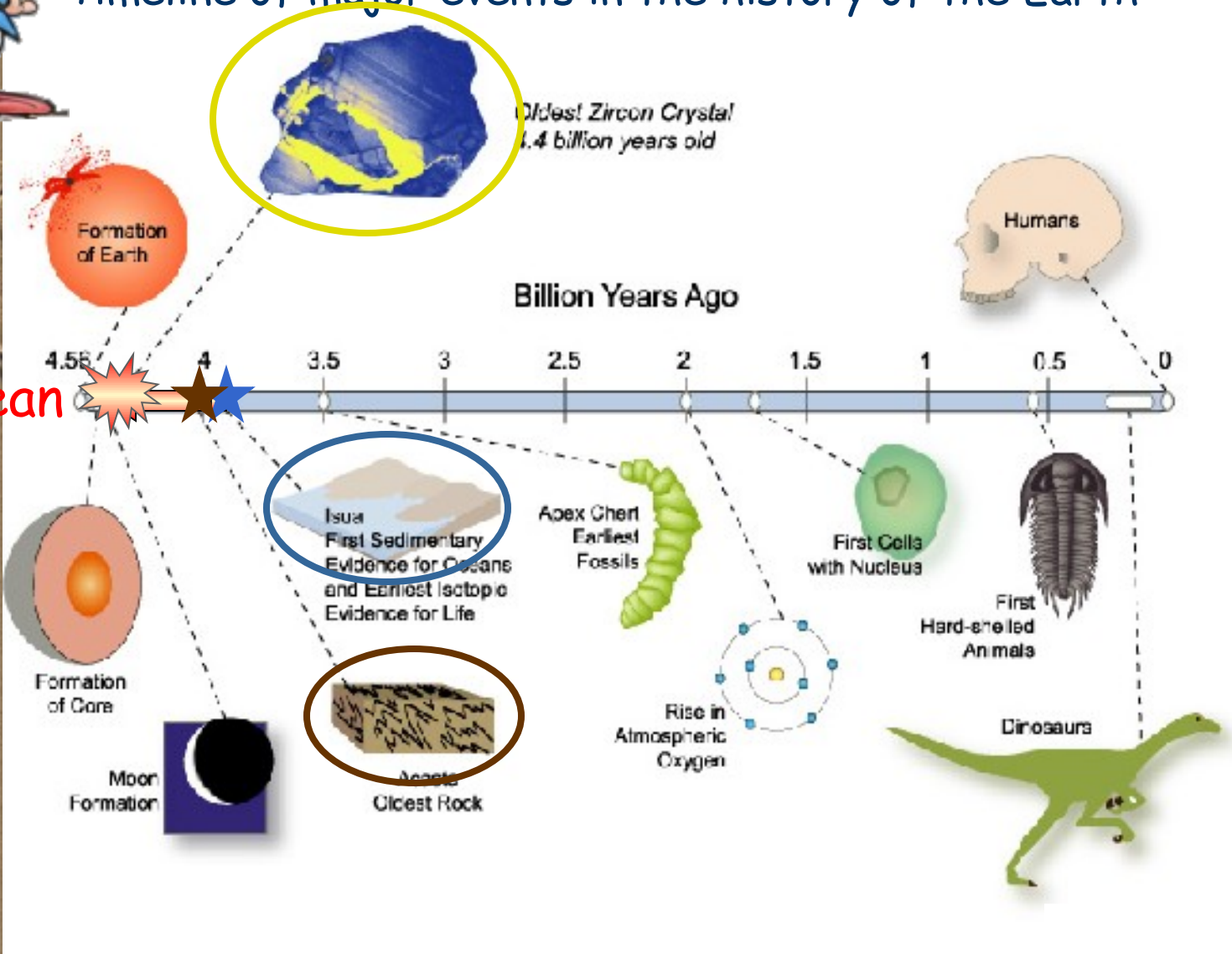
high
compatibility in zircon
slight

Continental crust growth models



How old is "old"?

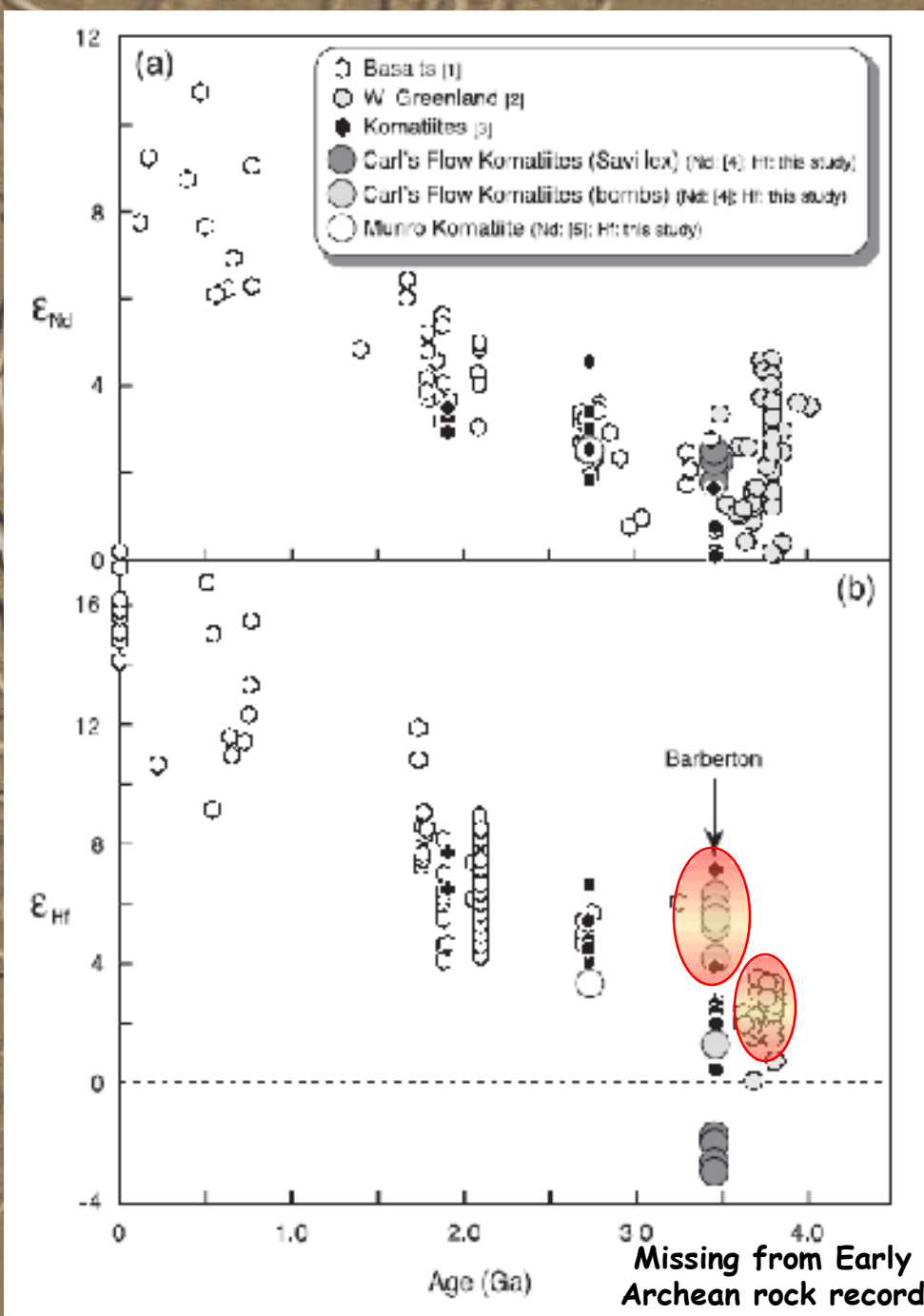
Timeline of major events in the history of the Earth



Hadean

$\epsilon_{Nd}(0) \sim +11$

$\epsilon_{Hf}(0) \sim +18$



Blichert-Toft et al. (1999, 2004)

Where is the Hadean crust?

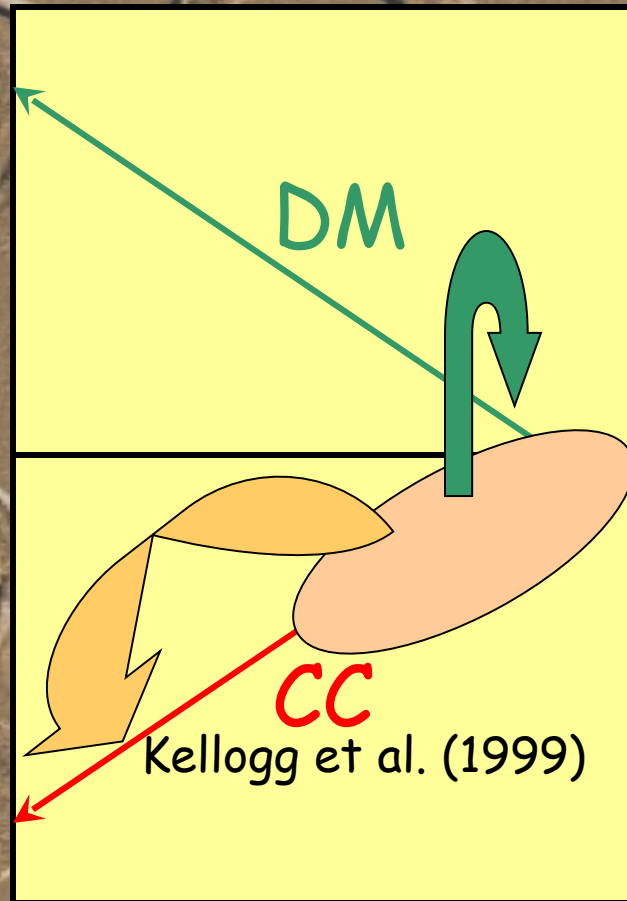
It never existed in the first place.....

Lost to space
by impacts



\$ =
0

Incorporated
into younger
crust?



Immediately
recycled
back into the
mantle?

Buried in the
so-called
"hidden"
reservoir?

Present

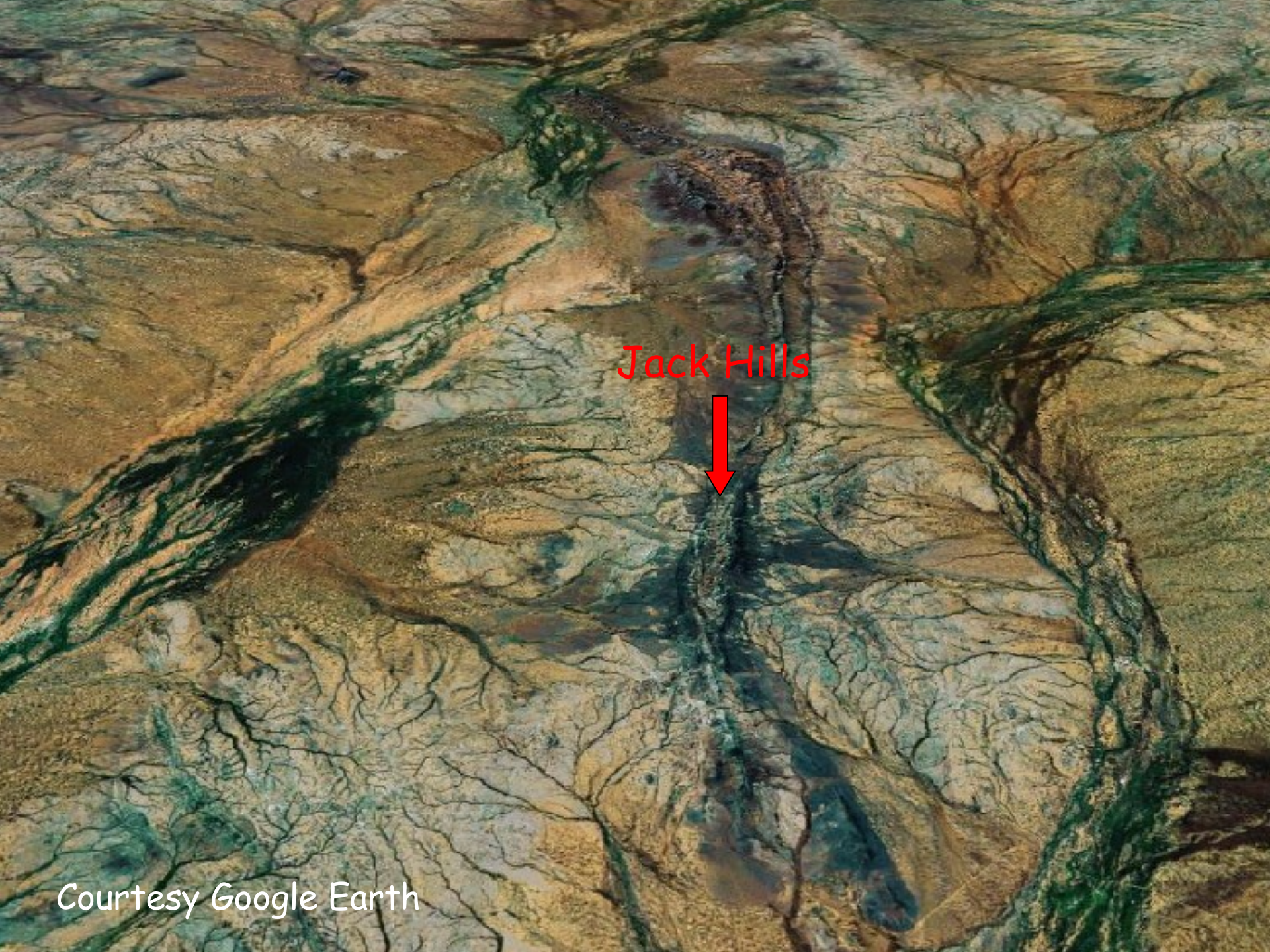
← Time (Ga)

4.0 4.5

Missing from Early Archean rock record

Les collines de Jack Hills, l'Australie de l'Ouest





Jack Hills



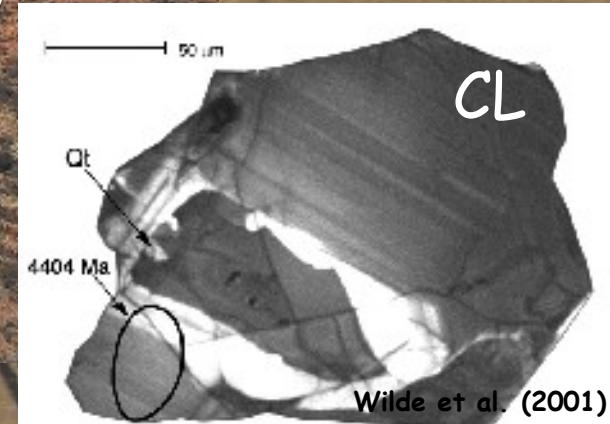
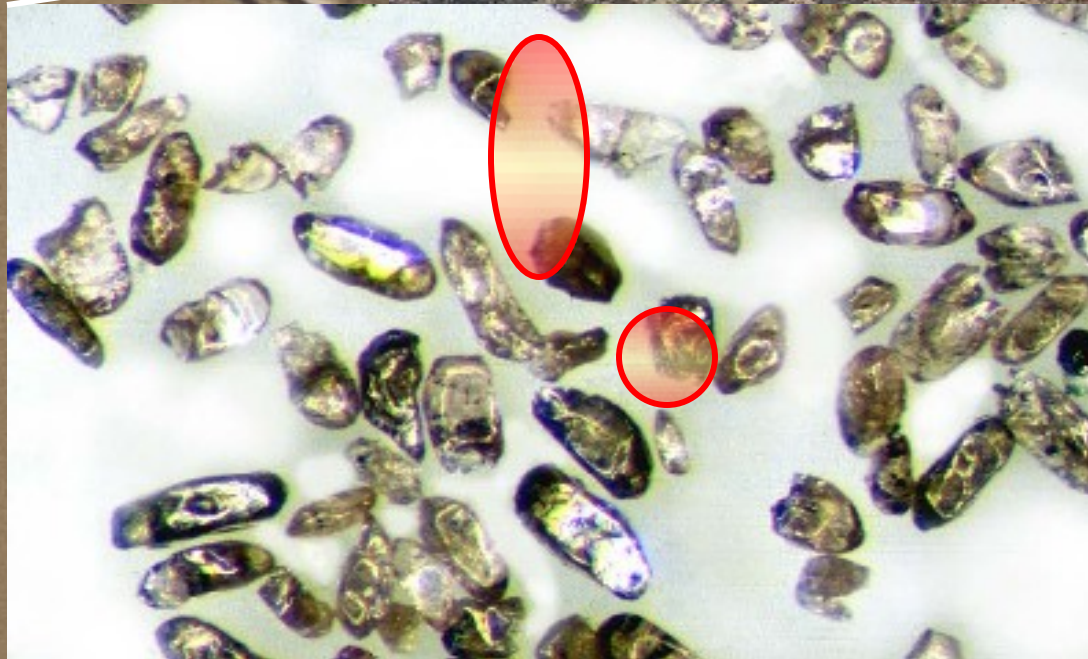
Courtesy Google Earth

A window on the missing years of Earth history

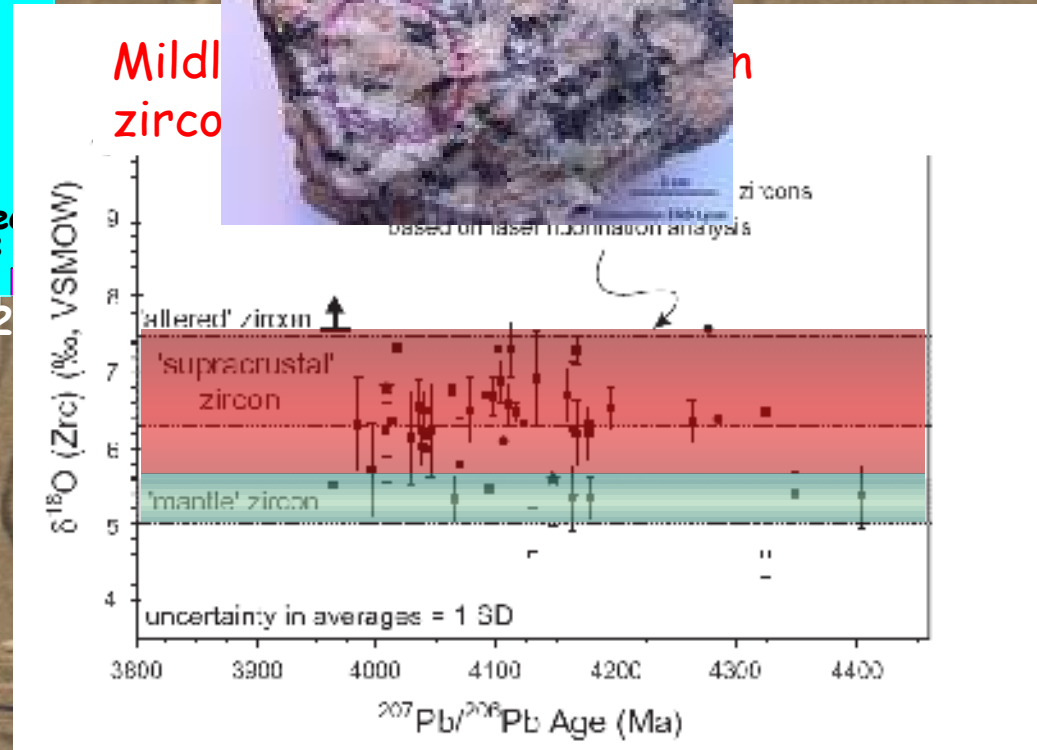
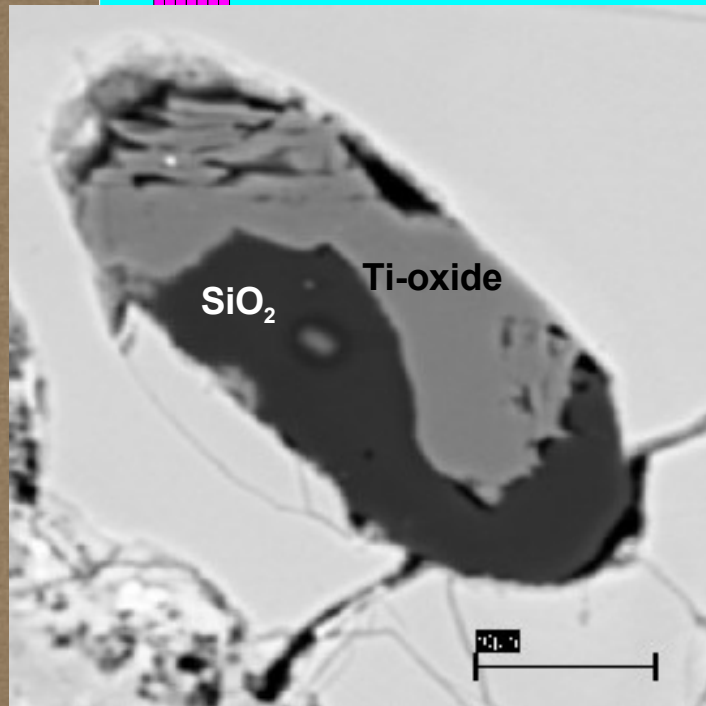
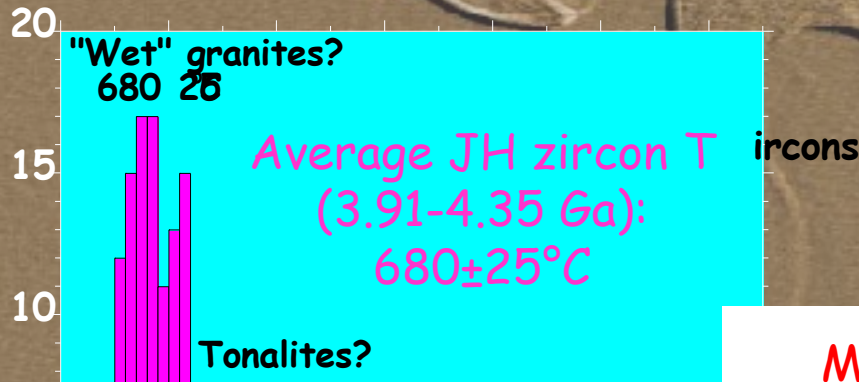
~3.1 Ga conglomerate



Ancient detrital zircon locality



What do we know about the host rock of the detrital Jack Hills zircons?

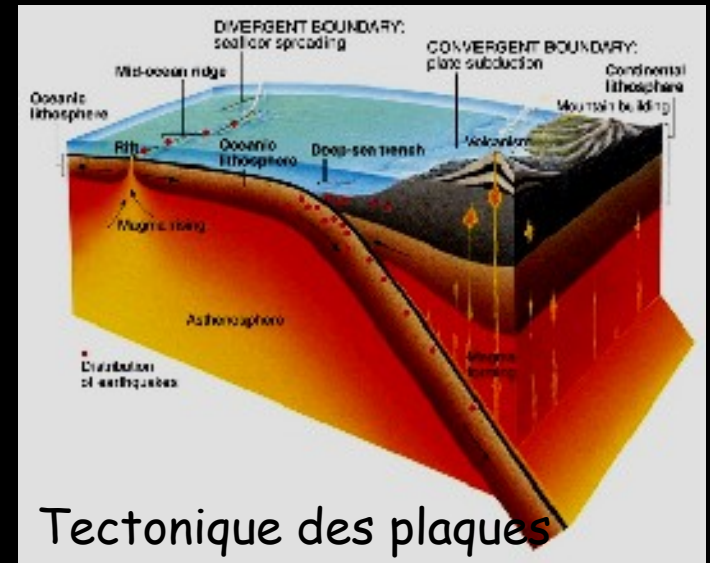


Pending issues:

- Multiplicity *vs* single event
 - Age of event(s)
- Nature, formation, and age of protocrust
- Implications for plate tectonics



Vie

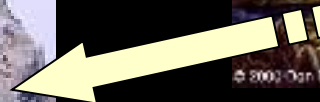


La géodynamique terrestre moderne
transporte des nutriments (NO_3 , PO_4 , SiO_2)
et accueille l'activité biologique.....

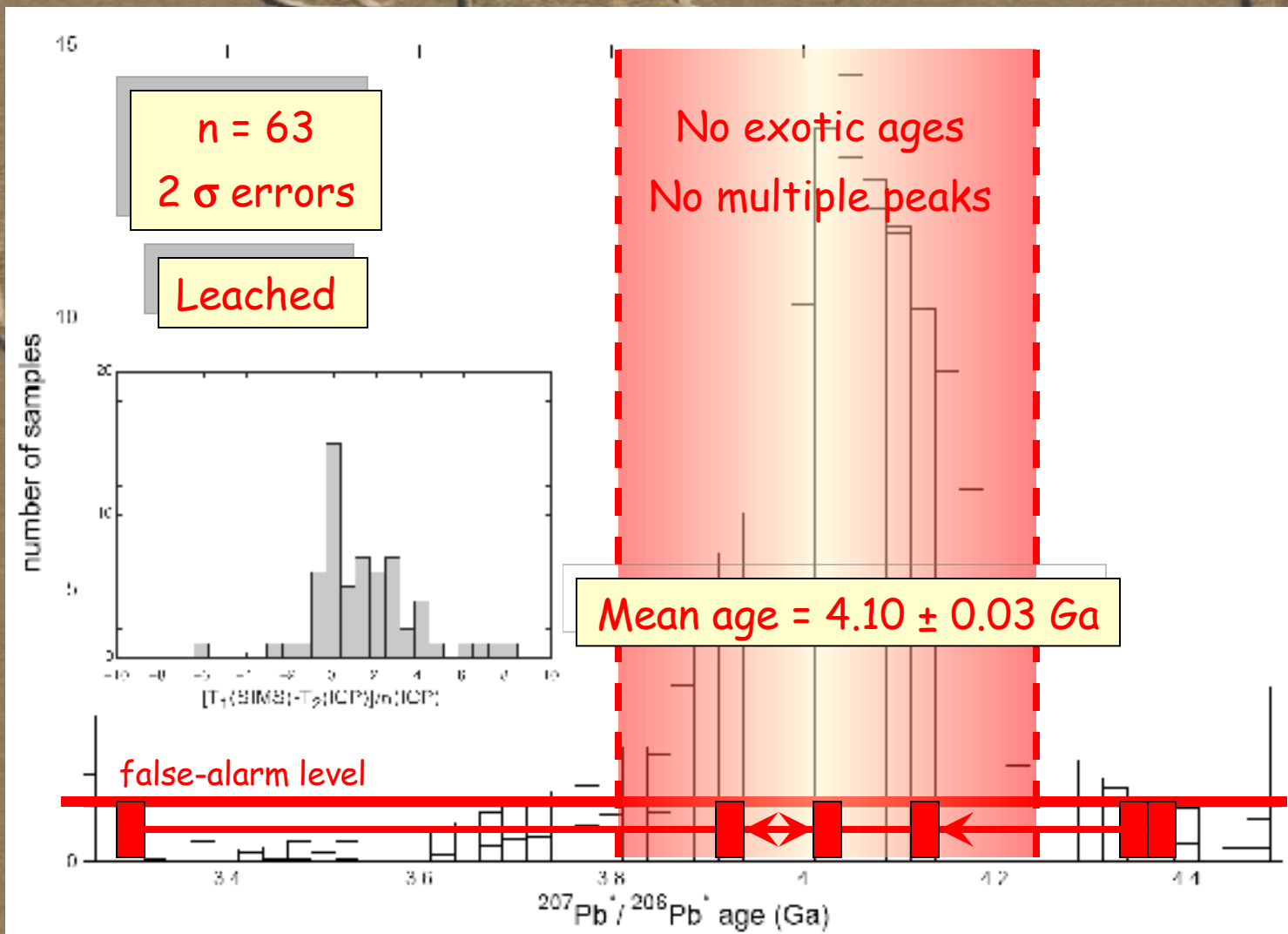
Quand?



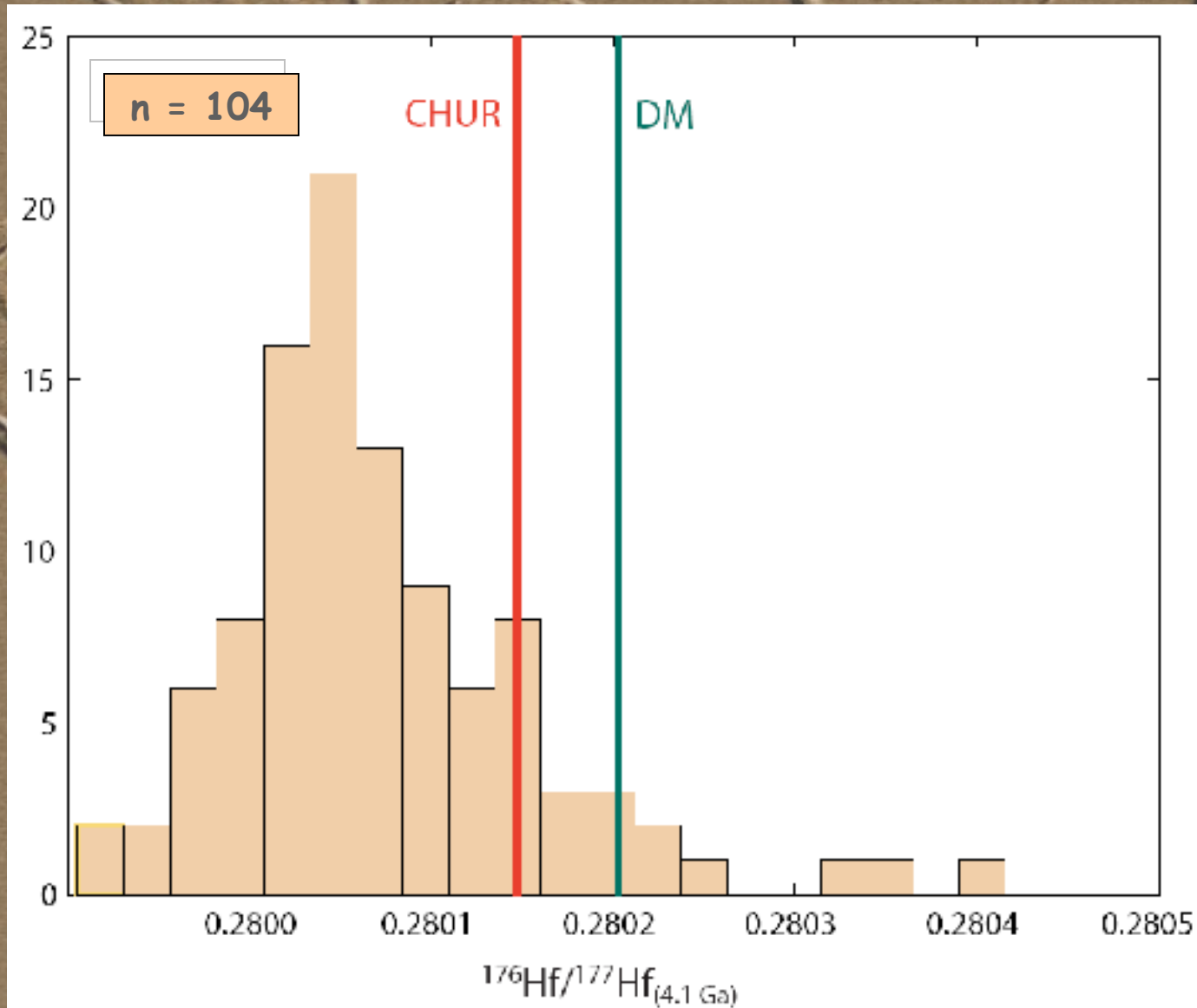
Granite



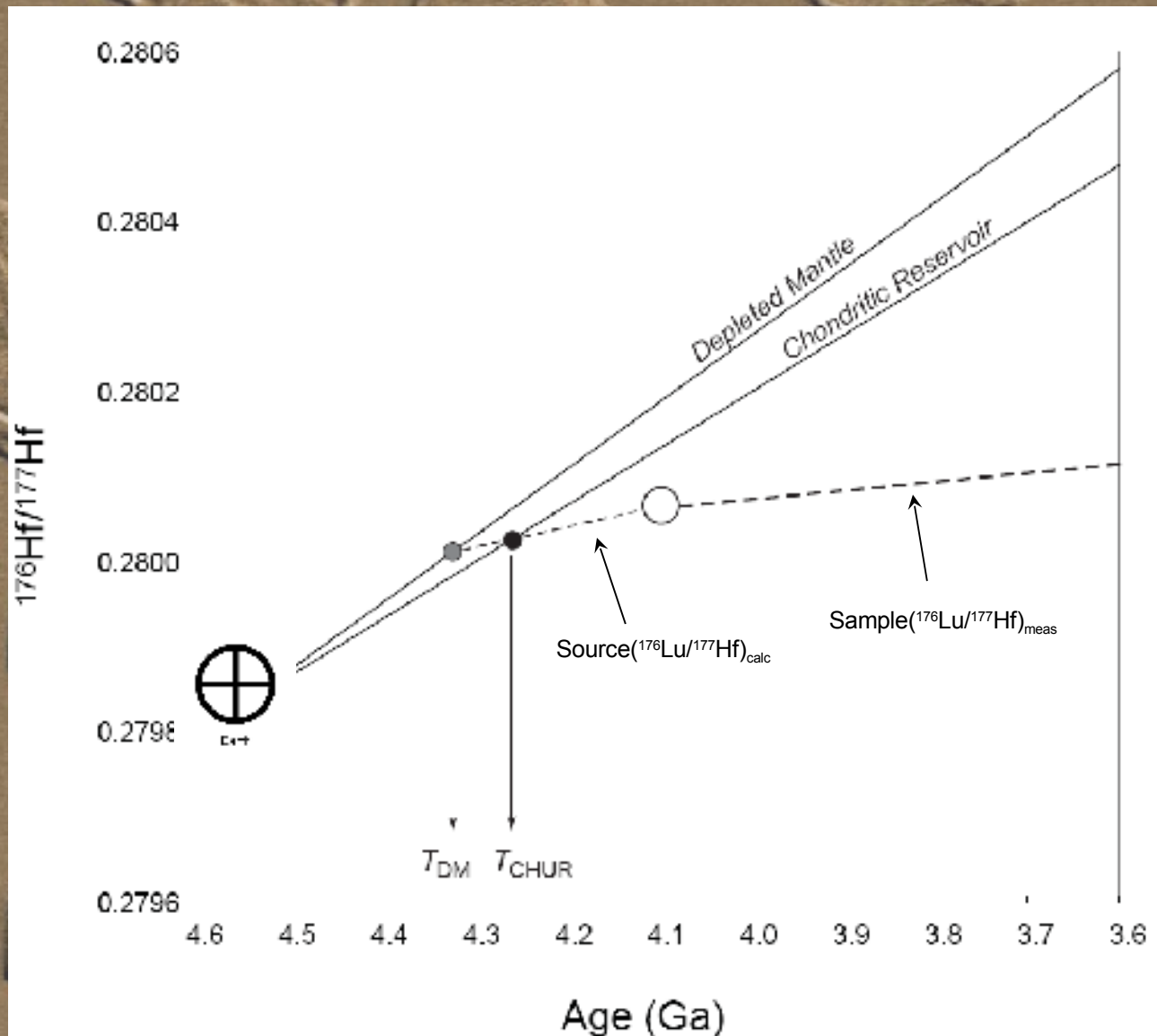
L'âge des plus vieux zircons terrestres = l'âge des granites de Jack Hills



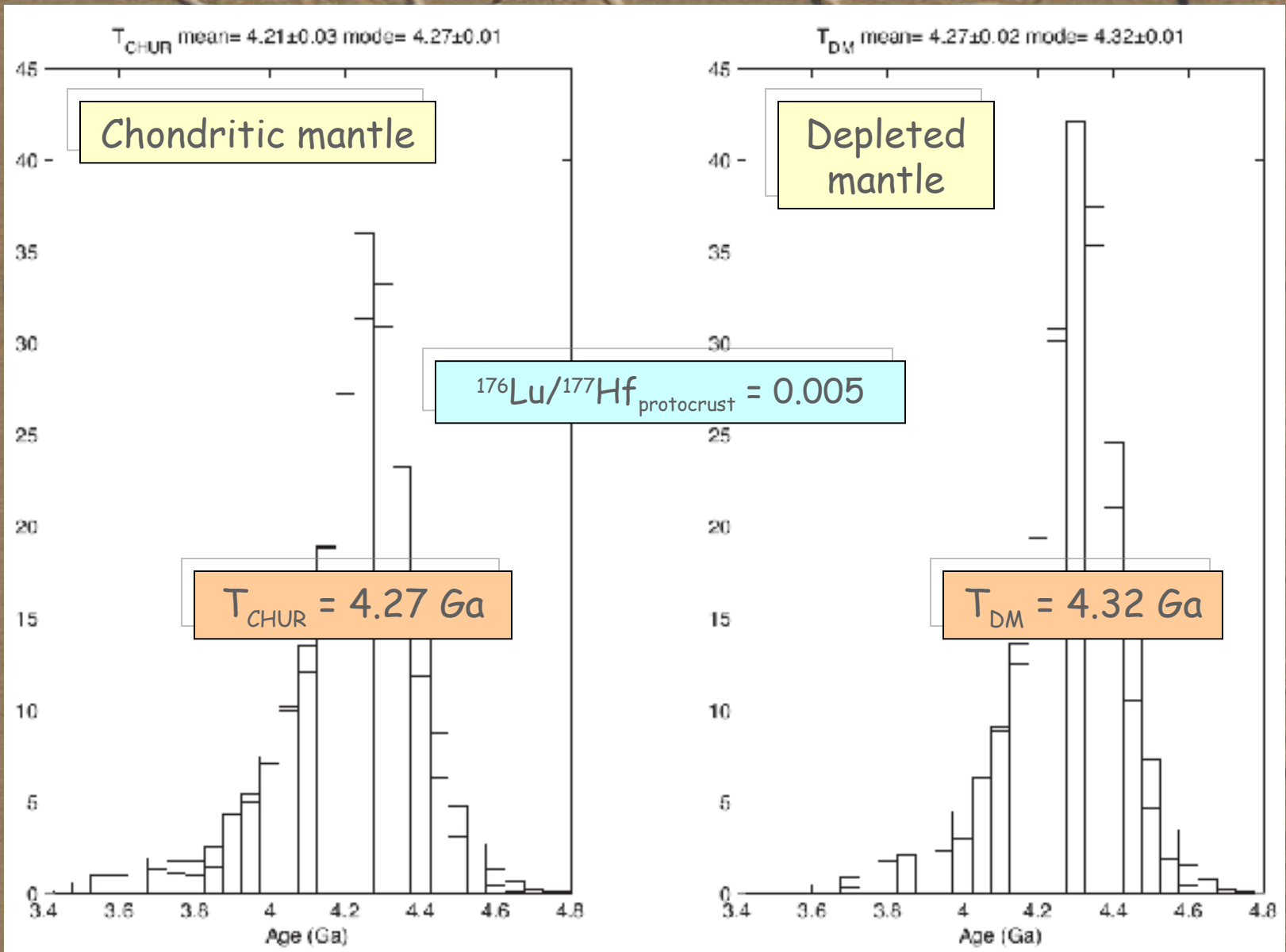
Initial Hf isotopic composition of Jack Hills zircons



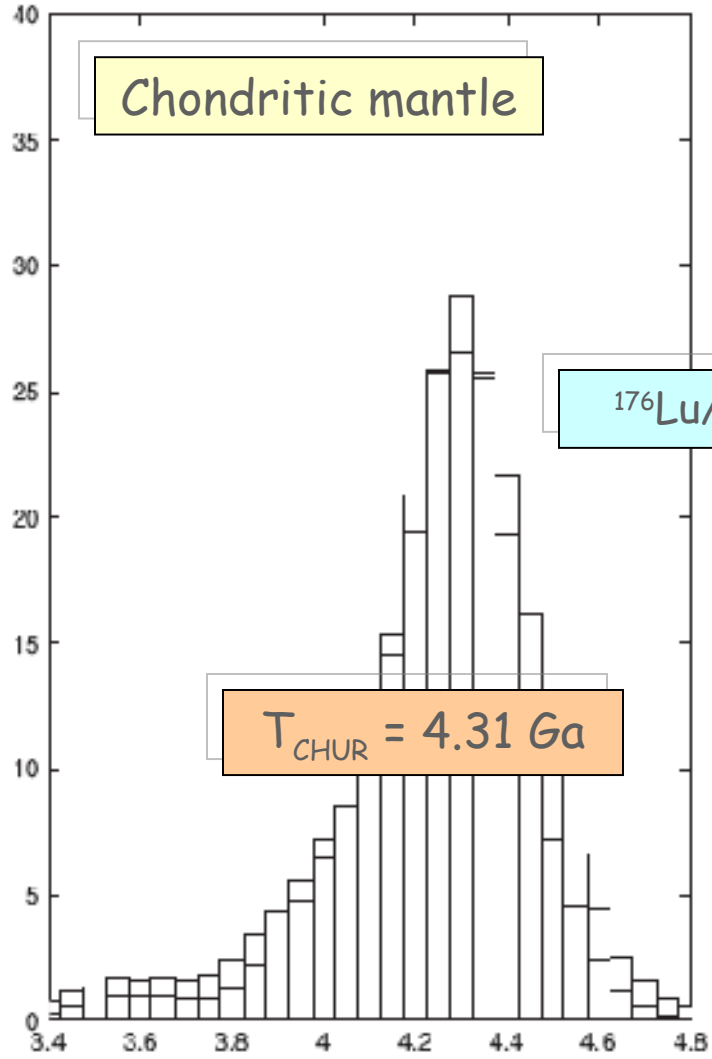
Rappel: Les âges modèles



L'âge ^{176}Lu - ^{176}Hf des plus vieux continents parents des granites de Jack Hills

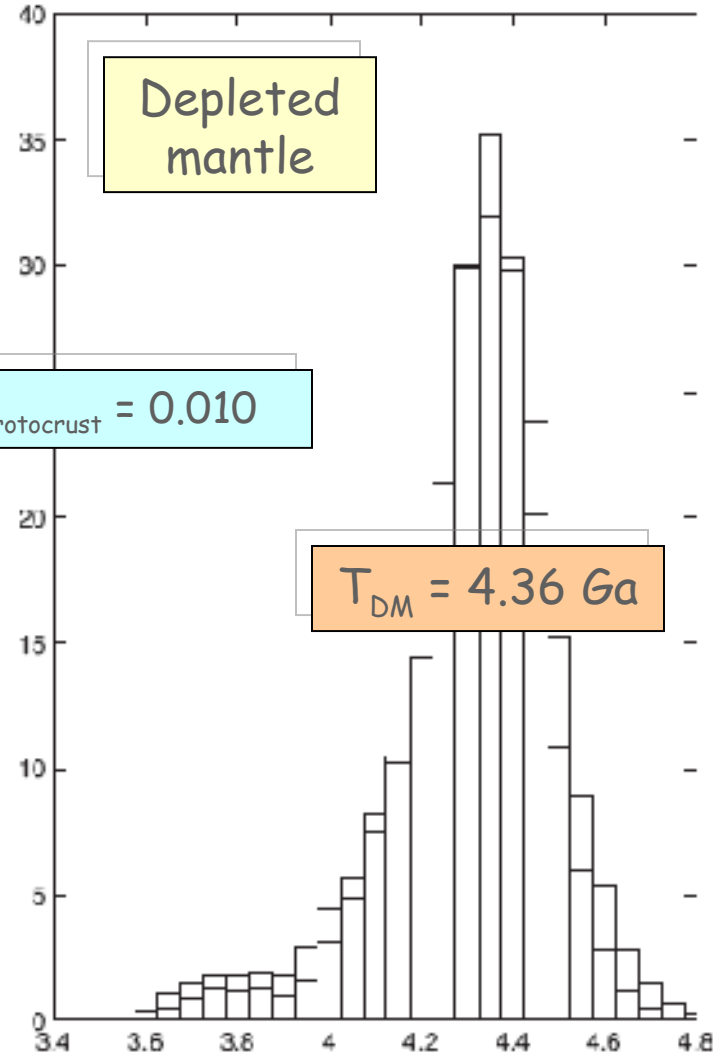


T_{CHUR} mean= 4.24 ± 0.04 Ga mode= 4.31 ± 0.01 Ga

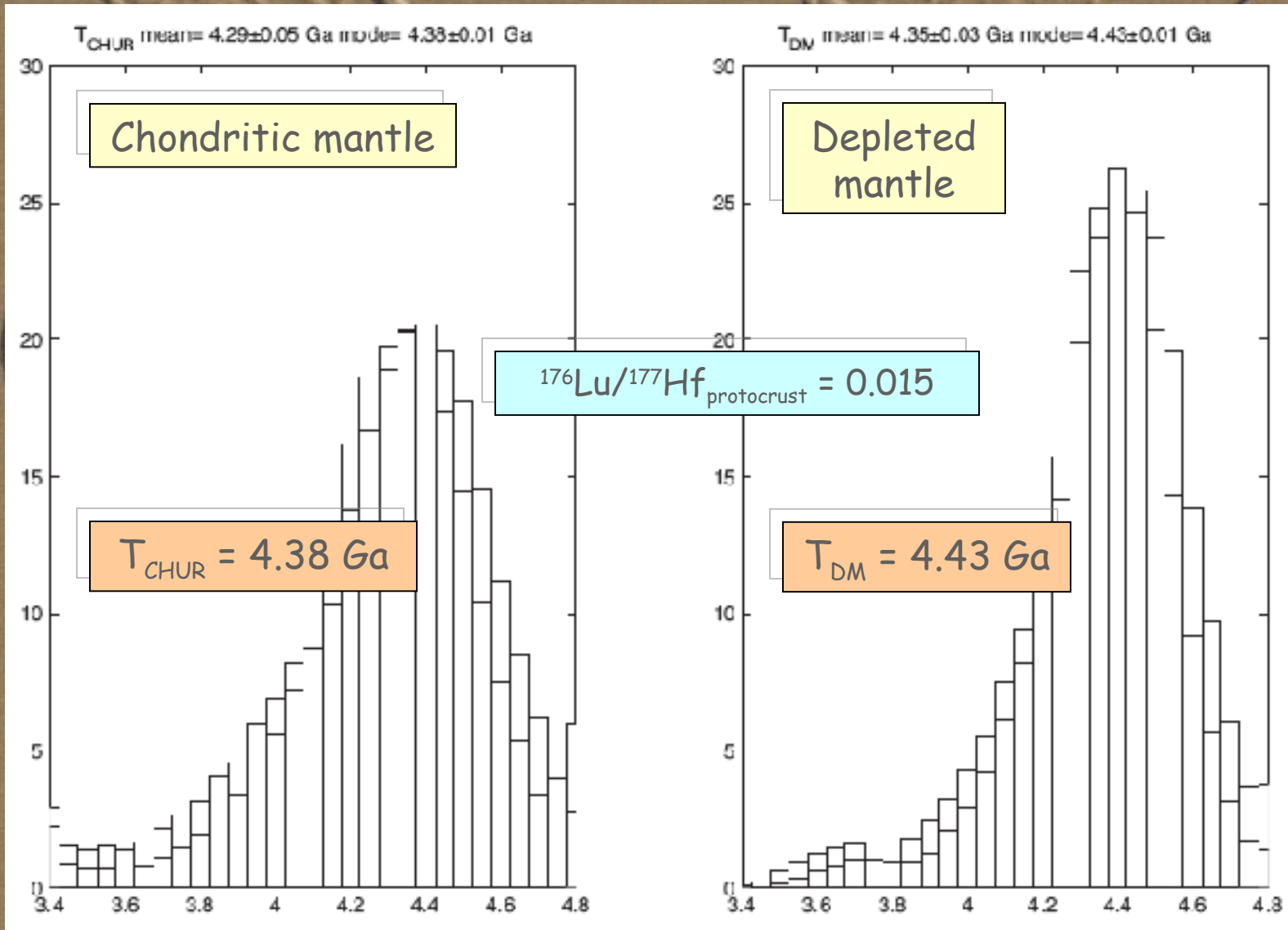


Age (Ga)

T_{DM} mean= 4.31 ± 0.03 Ga mode= 4.35 ± 0.01 Ga

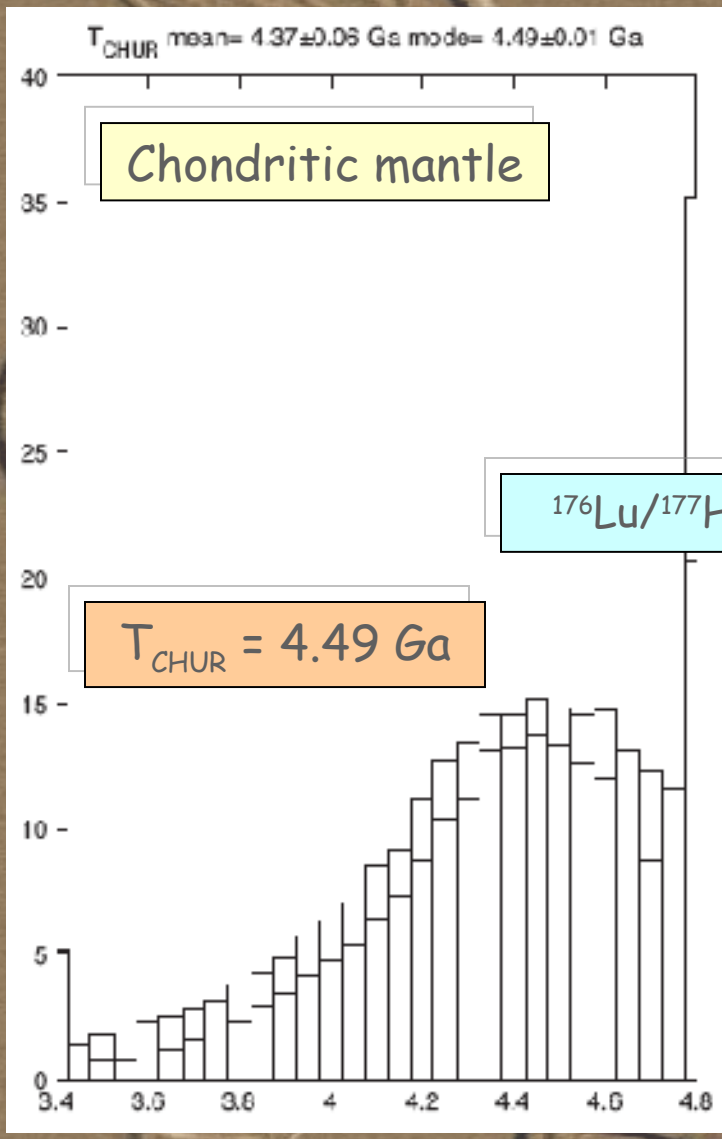


Age (Ga)

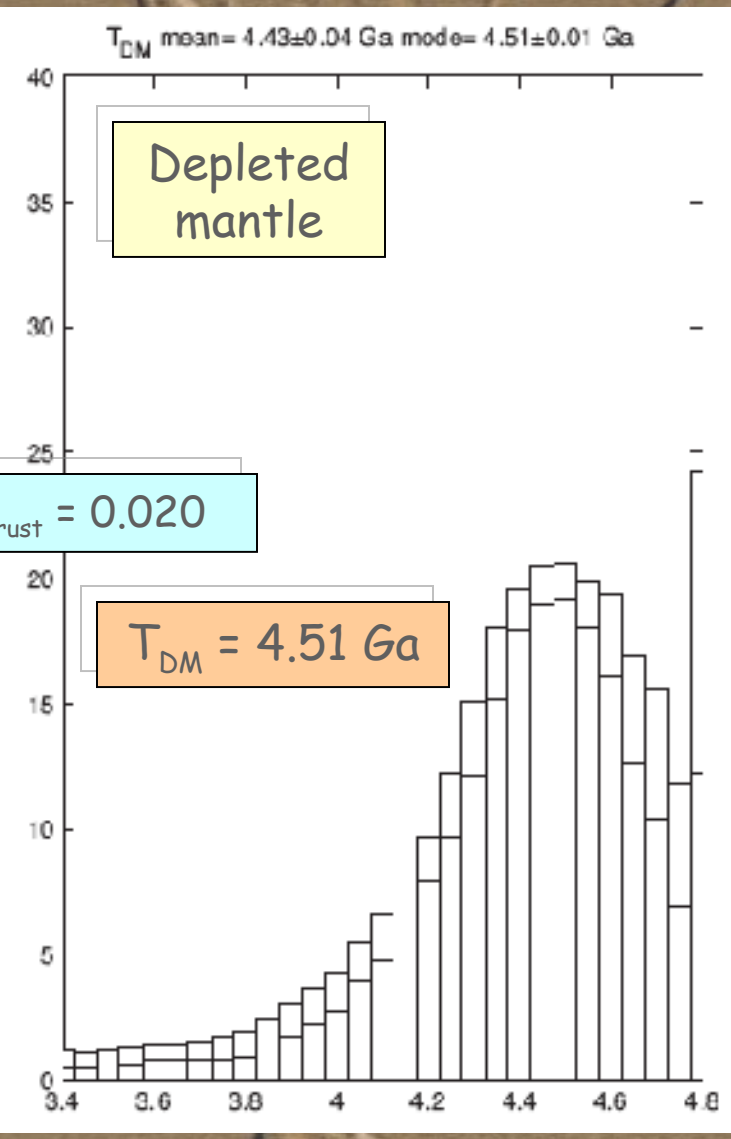


Age (Ga)

Age (Ga)

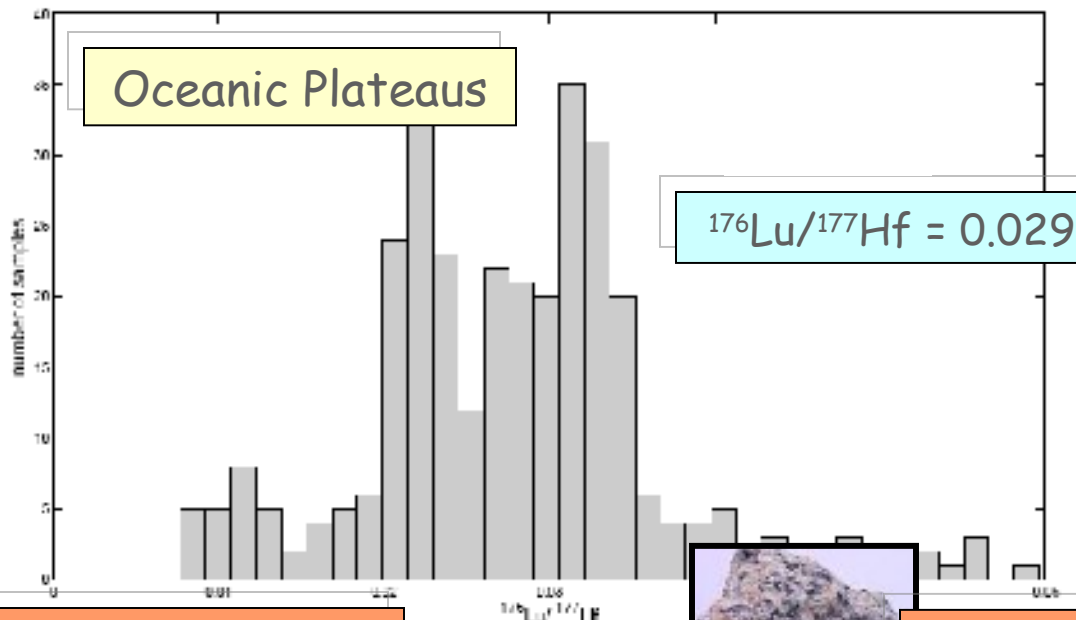


Age (Ga)

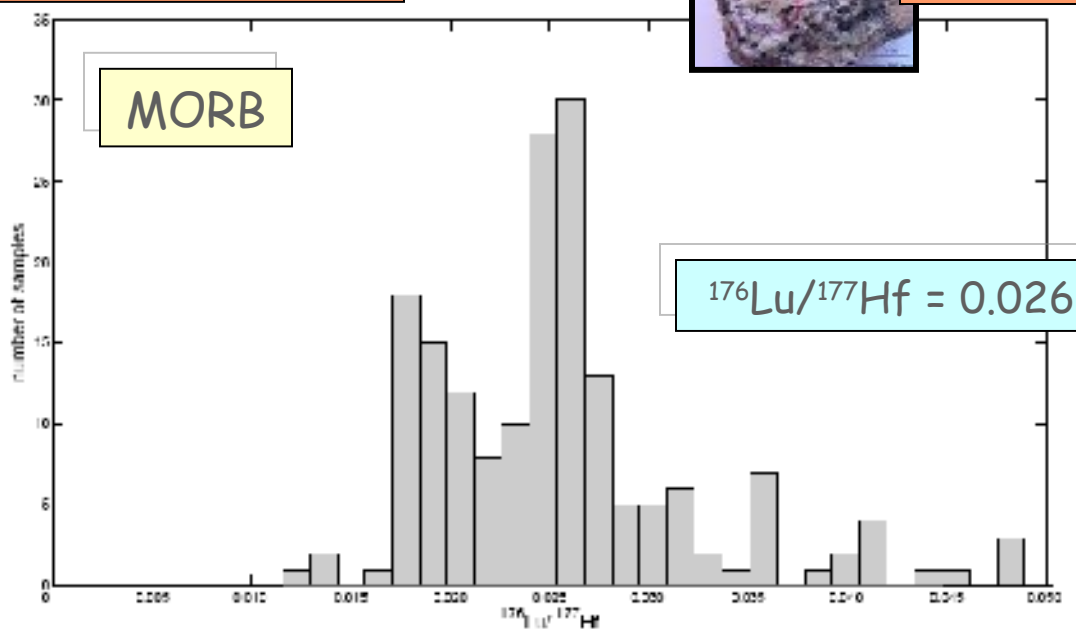


Age (Ga)

$^{176}\text{Lu}/^{177}\text{Hf}_{\text{protocrust}} = 0.020$



$^{176}\text{Lu}/^{177}\text{Hf}_{\text{CC}} = 0.012$



First-order observations:

- $^{176}\text{Lu}/^{177}\text{Hf} \geq 0.015$, as in crust made of basalt, produces flat histograms with many unacceptable ages $> 4.56 \text{ Ga}$
- 'Granitic' $^{176}\text{Lu}/^{177}\text{Hf} = 0.005\text{-}0.010$ produces well-defined peaks centered around maxima at $4.31 (T_{\text{CHUR}})$ and $4.36 (T_{\text{DM}}) \text{ Ga}$ with all ages $< 4.56 \text{ Ga}$



Les premiers continents apparaissent alors
~200 Ma après la formation de la Terre

Pending issues:

- Multiplicity vs single event
 - Age of event(s)
- Nature, formation, and of protocrust age
- Implications for plate tectonics

Because granites are the hallmark of modern plate tectonics, the Jack Hills zircon bulk Pb ages suggest this process was active in its present form by at least 4.1 Ga

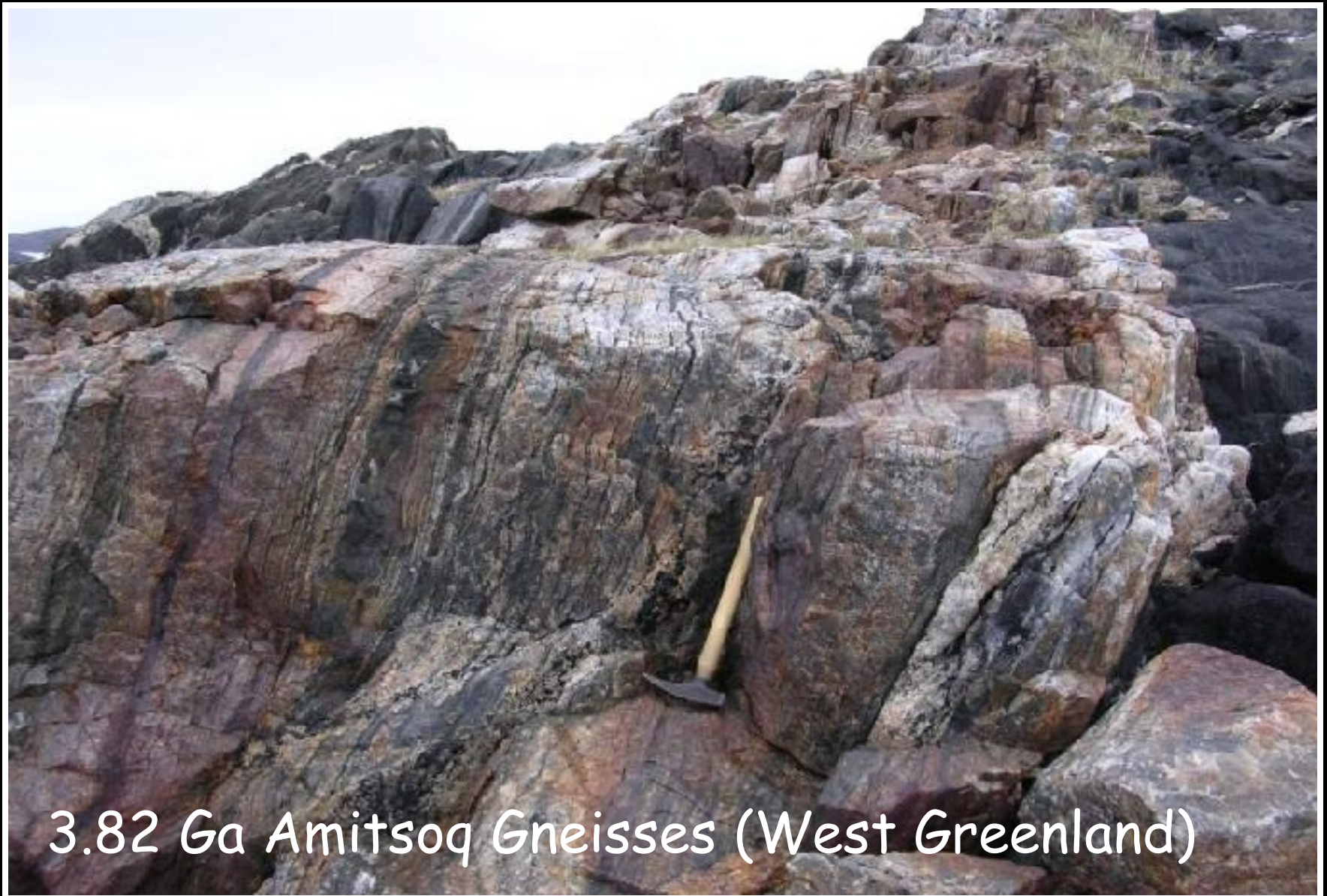
Demonstrating it started even earlier requires independent evidence that the ~4.35 Ga source rock of the Jack Hills granites was *also* granitic

This *could* be given the

- heavy oxygen
- unradiogenic Hf
- low Lu/Hf

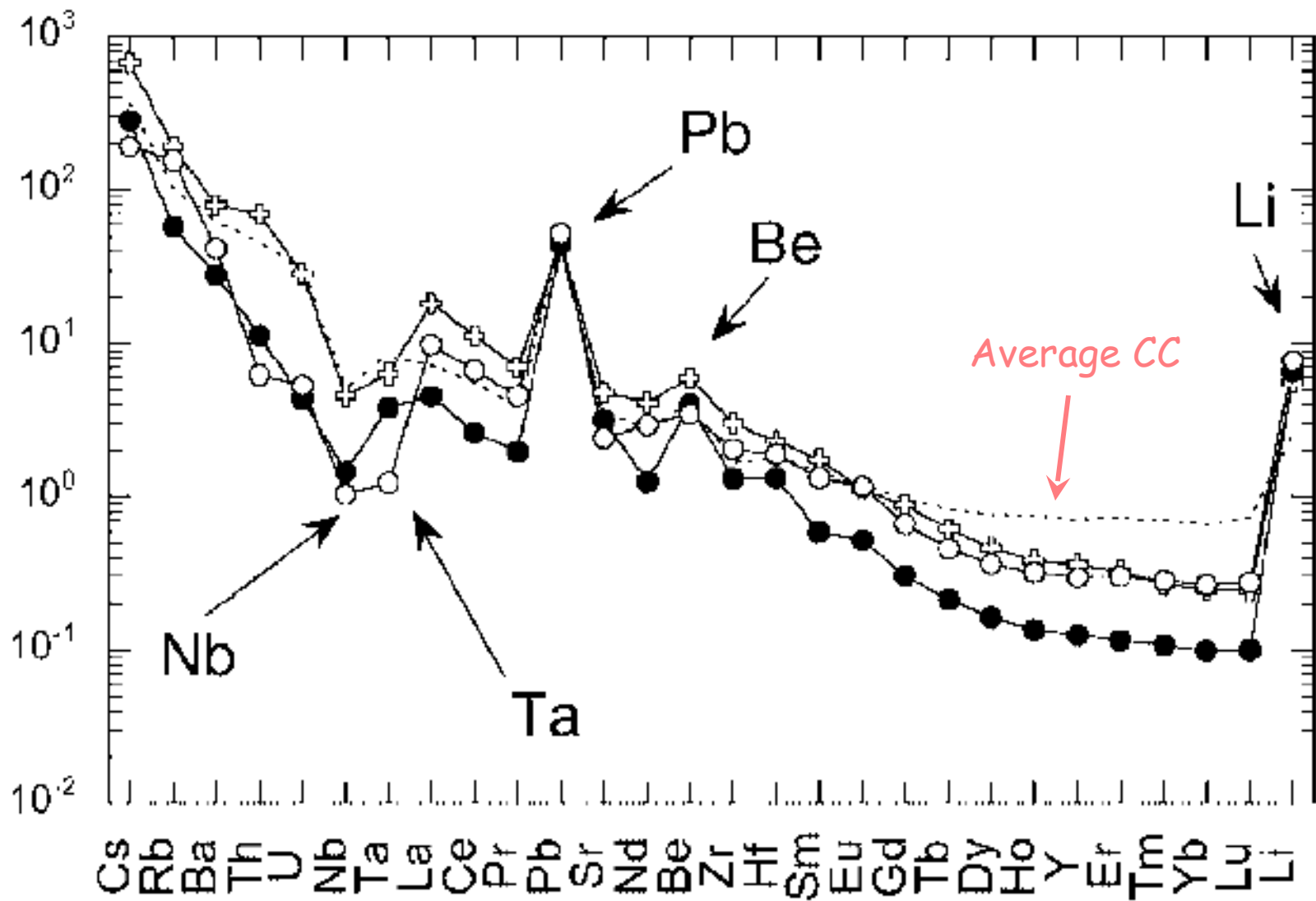
BUT.....Lu/Hf way too low!!!

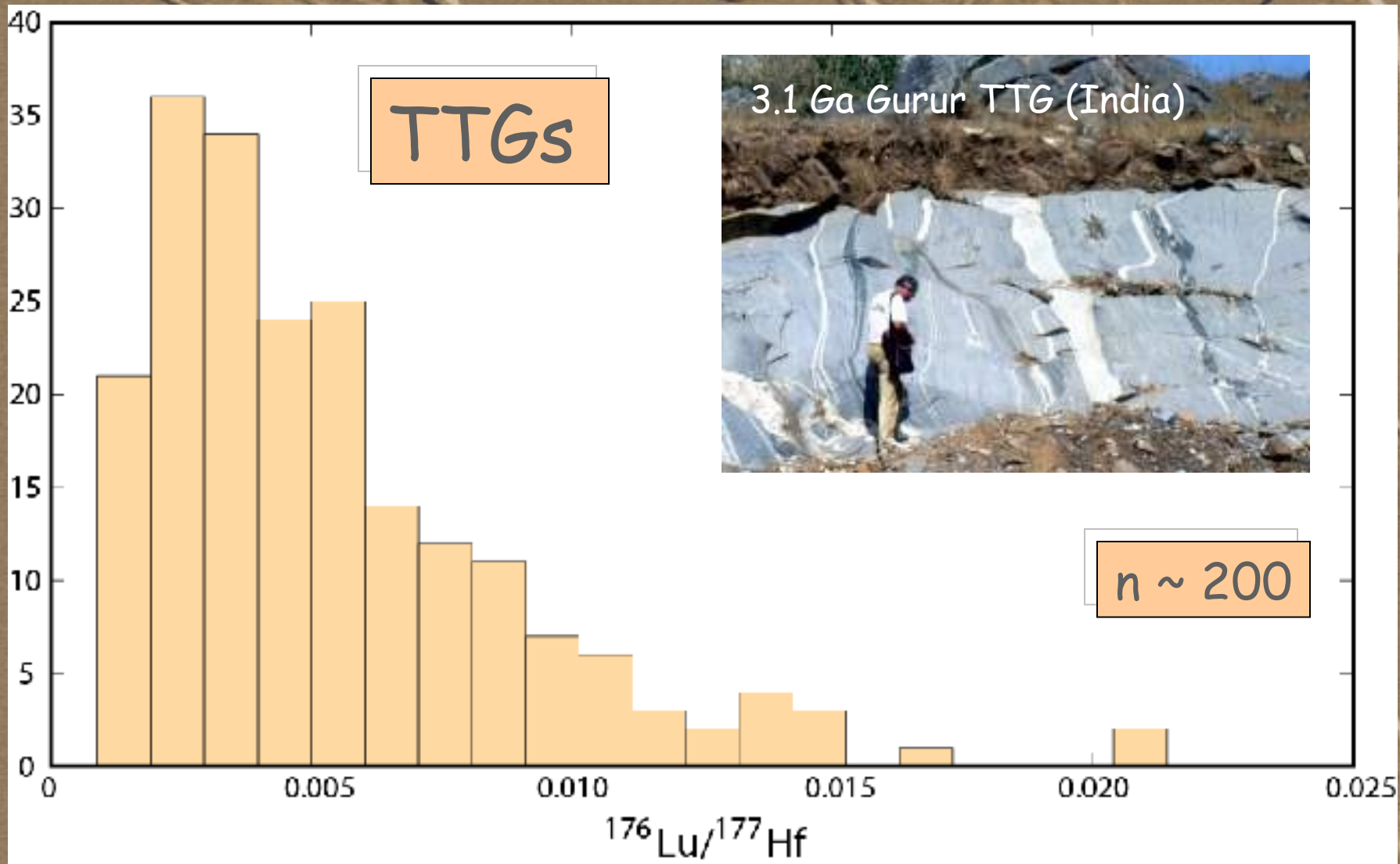
Only Archean TTGs have such inordinately low Lu/Hf



3.82 Ga Amitsoq Gneisses (West Greenland)

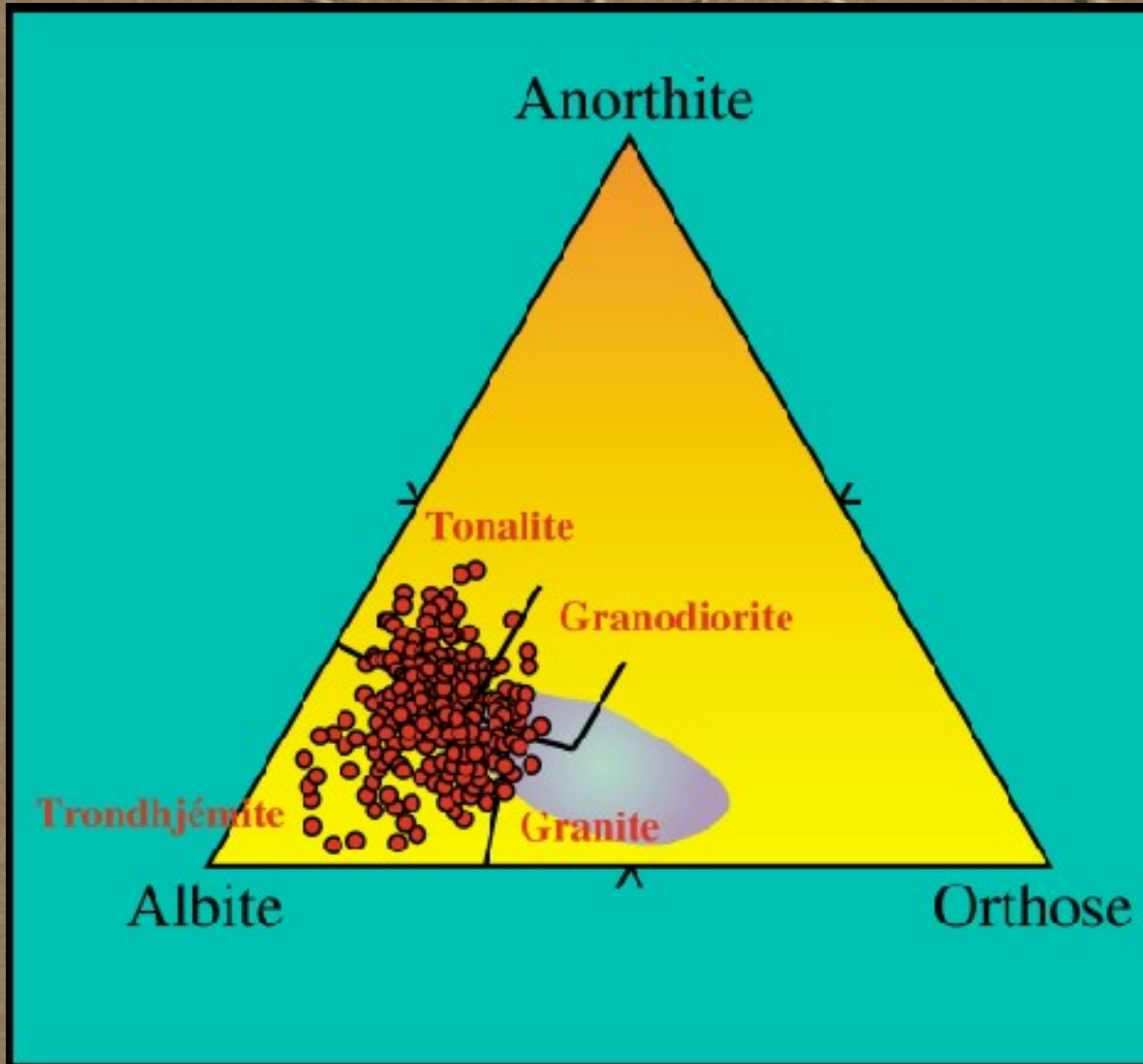
N-MORB normalised concentration





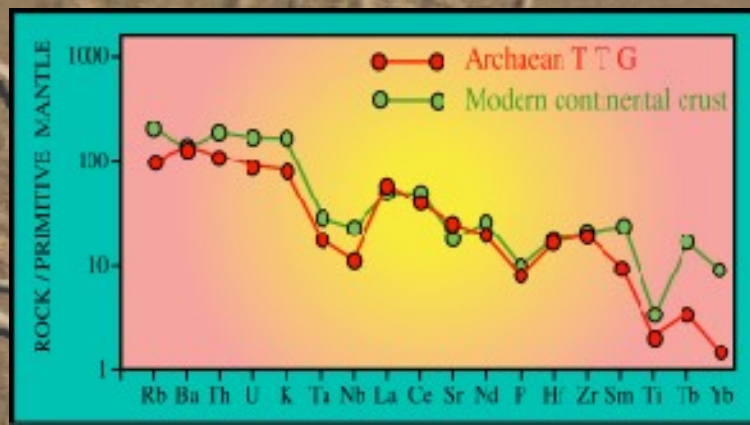
Hervé Martin (unpublished TTG data compilation)

Primitive continental crust



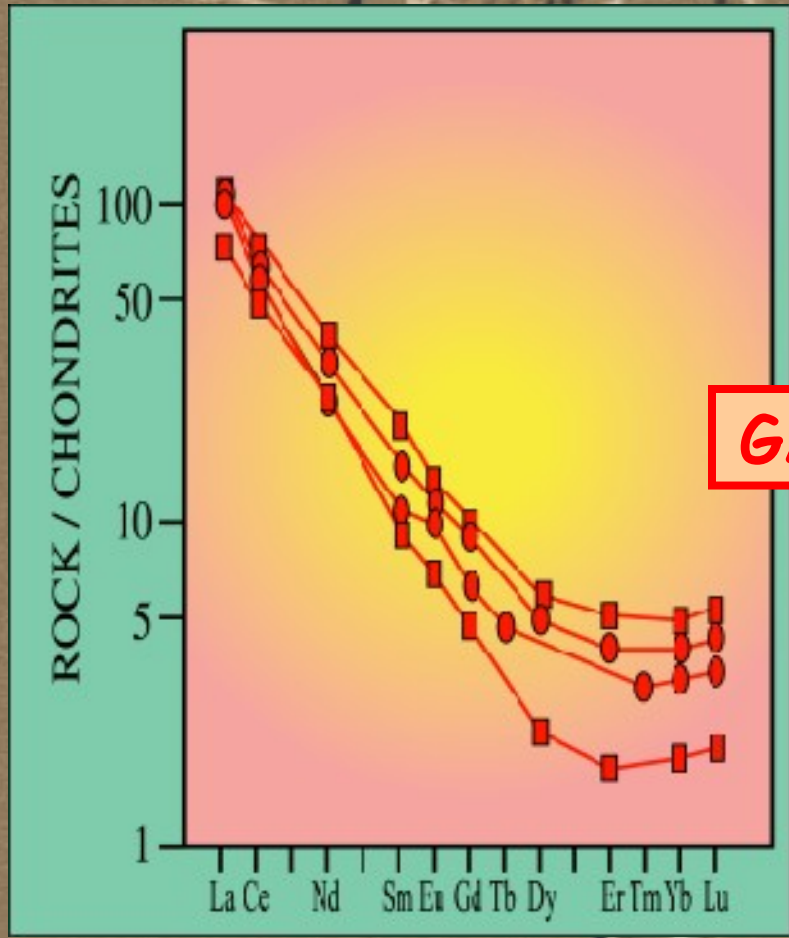
TTGs are calc-alkaline
but Na-rich:
Tonalitic
Trondhjemitic
Granodioritic

Modern continental
crust is "classic" calc-
alkaline = K-rich:
Granodioritic
to Granitic

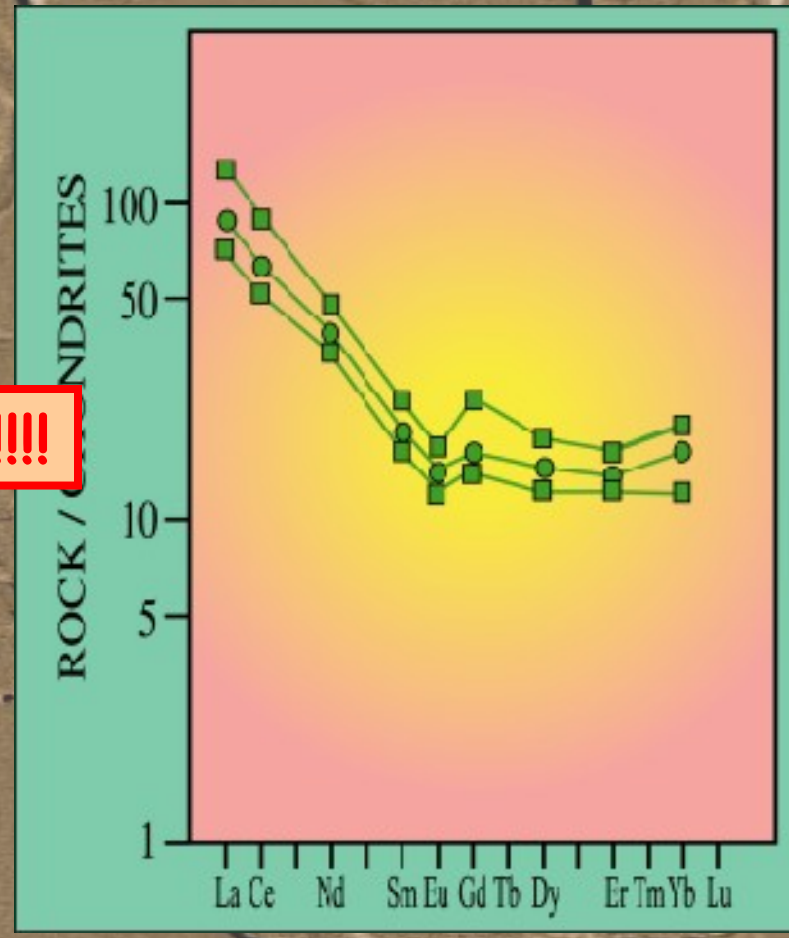


TTGs

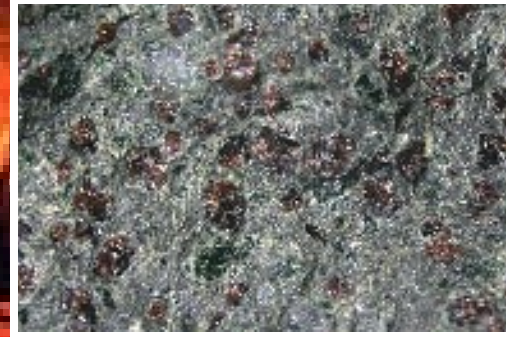
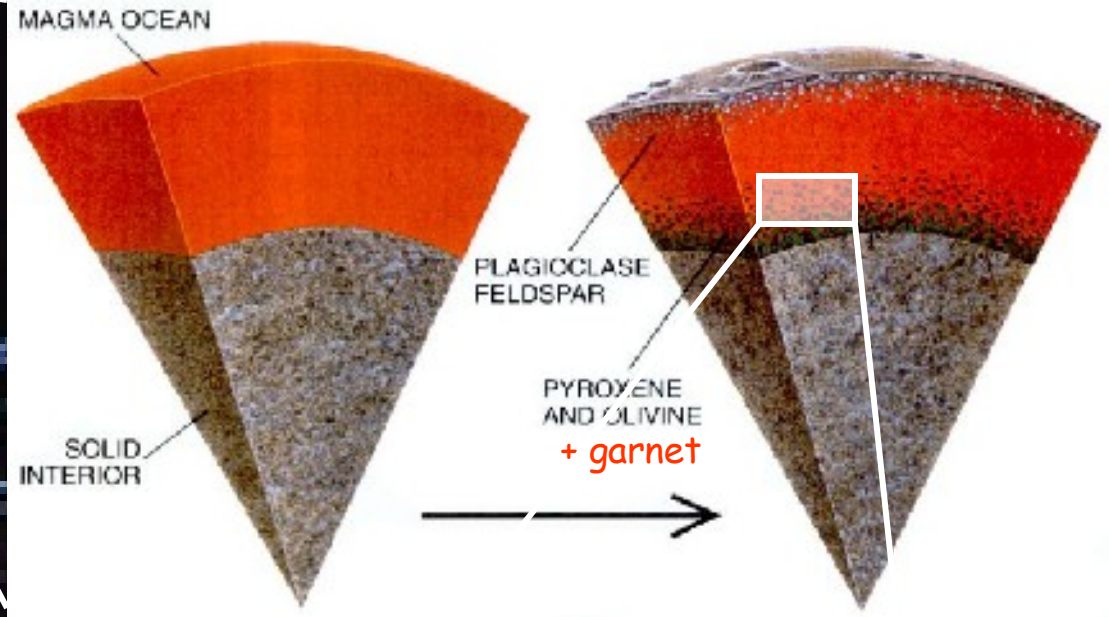
MODERN CC



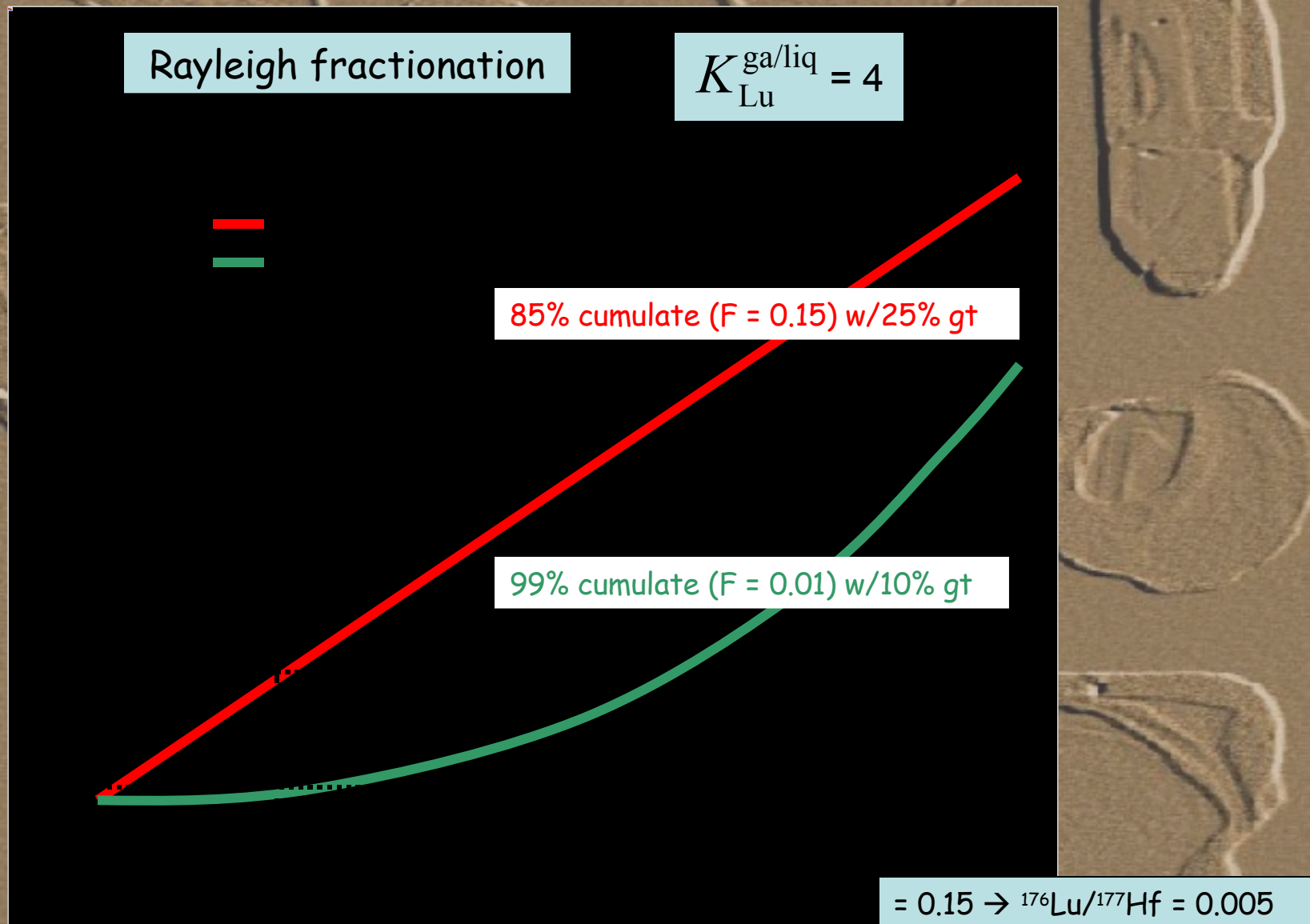
GARNET!!!!



Last stages of magma ocean crystallization



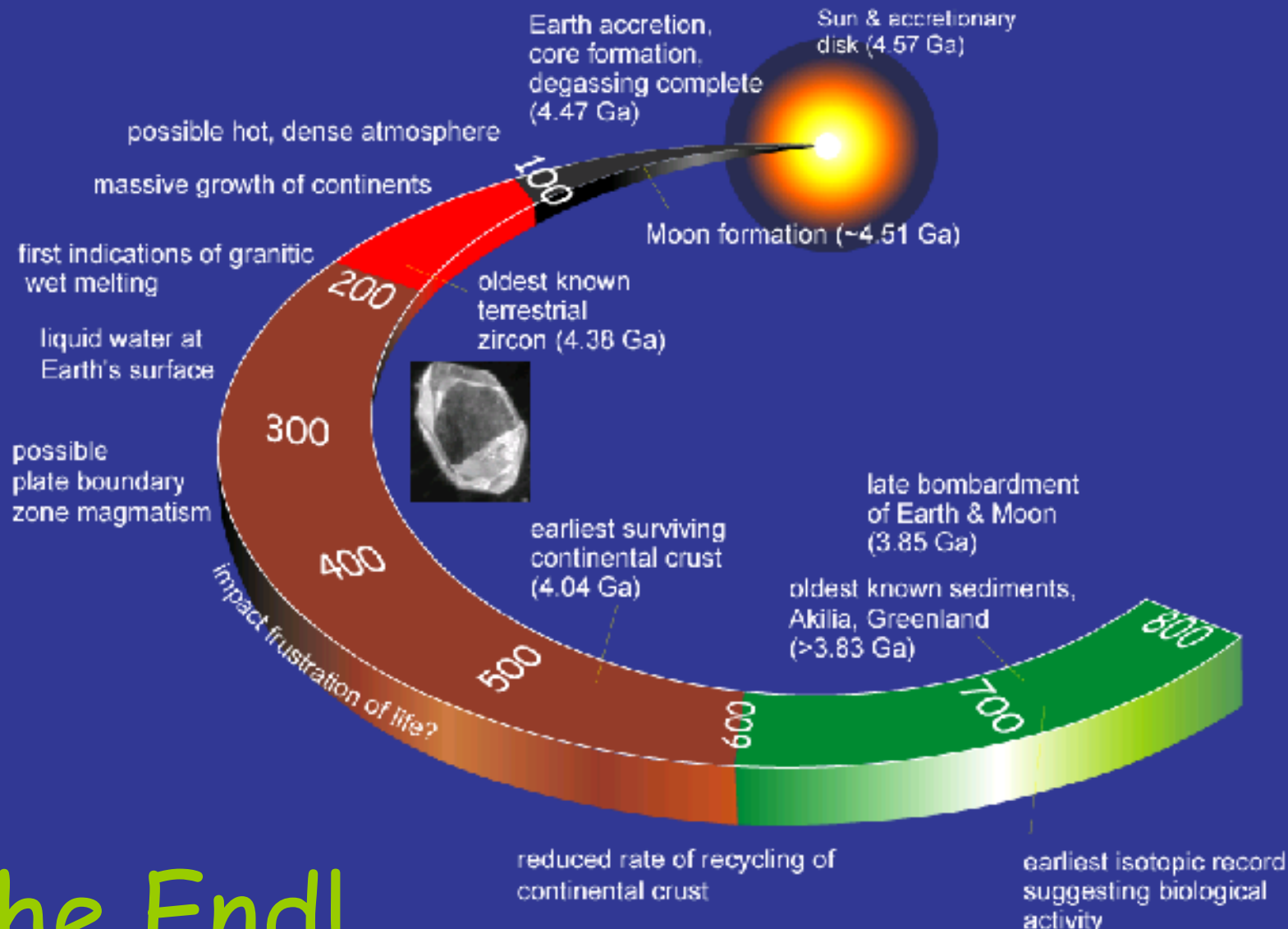
How did the early low-Lu/Hf crust form?



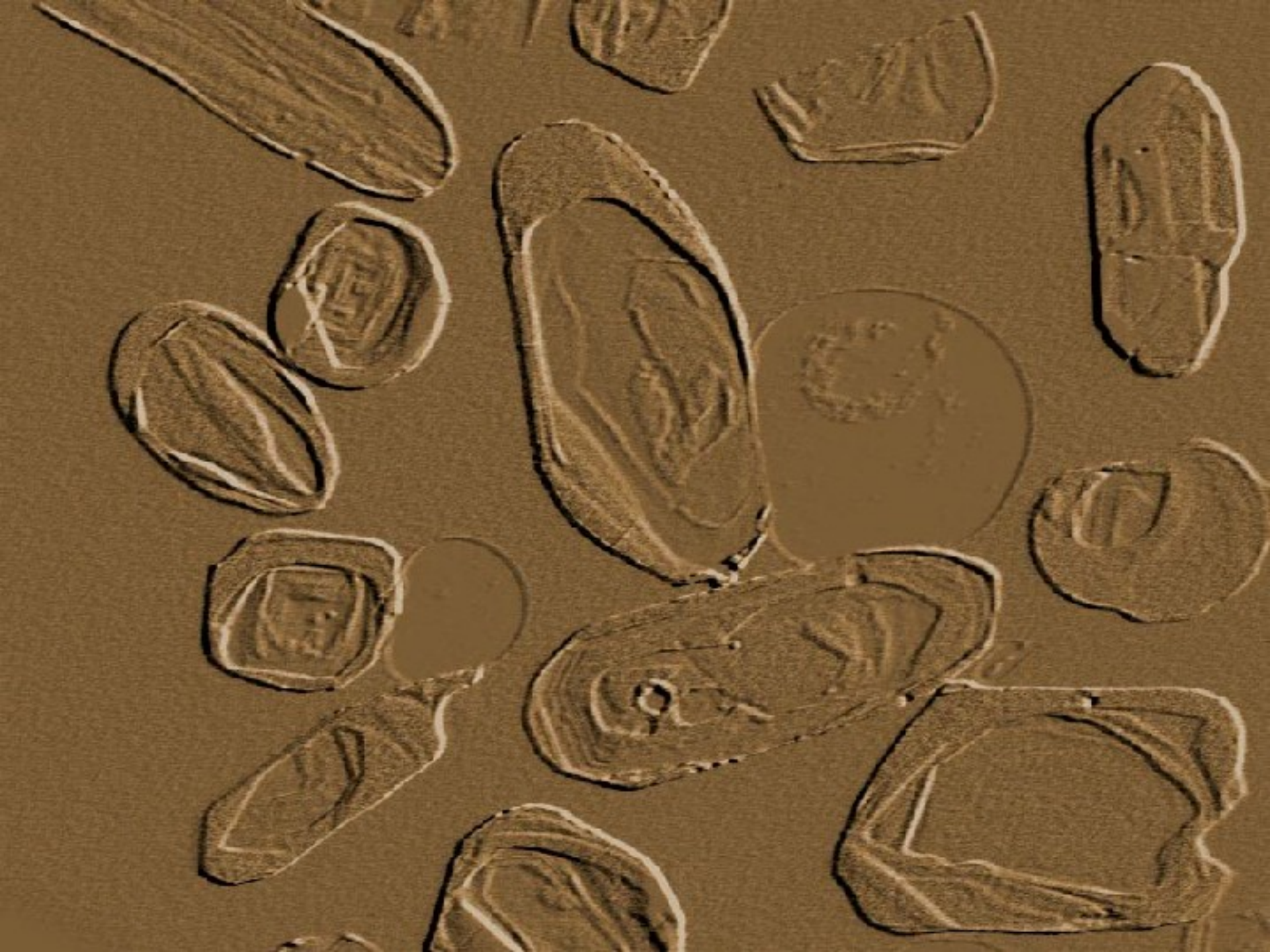
Conclusions

- The parent granites of the Jack Hills zircons formed 4.1 ± 0.1 Gy ago by remelting of a 4.31-4.36 Ga old protocrust...
- ...whose inferred low Lu/Hf ratio best fits Archean TTG suites...
- ... which themselves may have been produced by melting of the last remains of the magma ocean (hydrous KREEPy basalt)
- Dating the onset of plate tectonics hinges around whether TTGs are subduction zone magmas or not

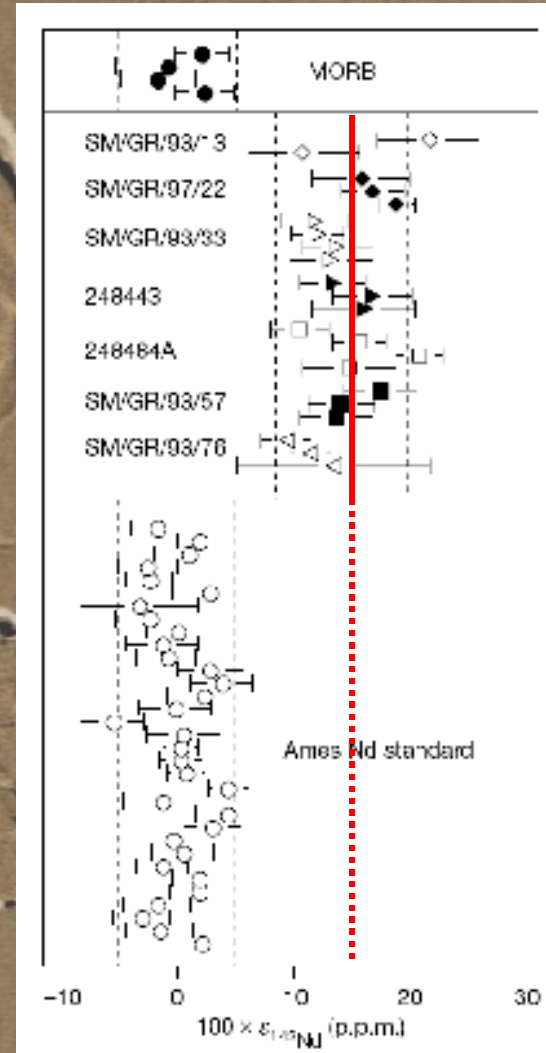
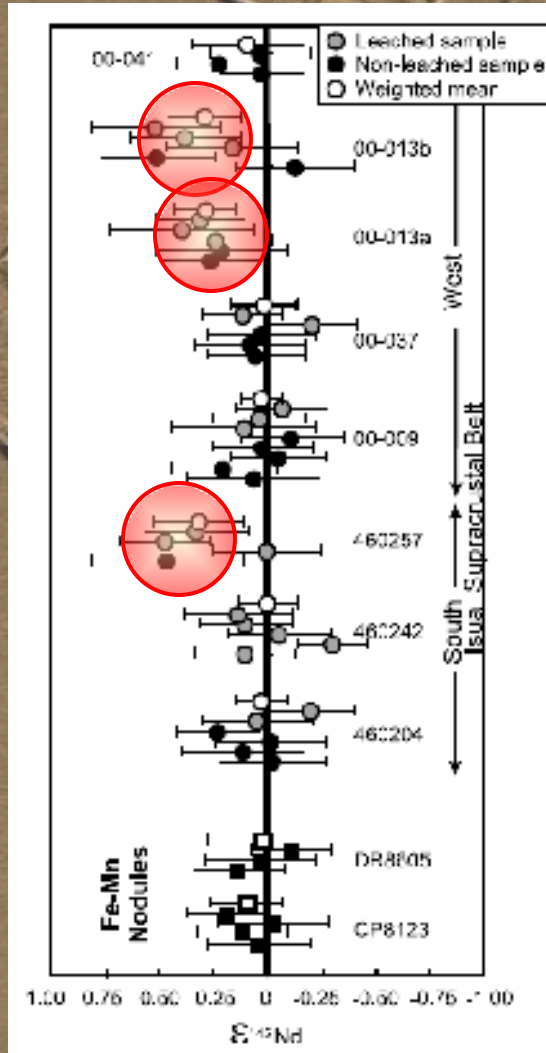
The first 800 million years



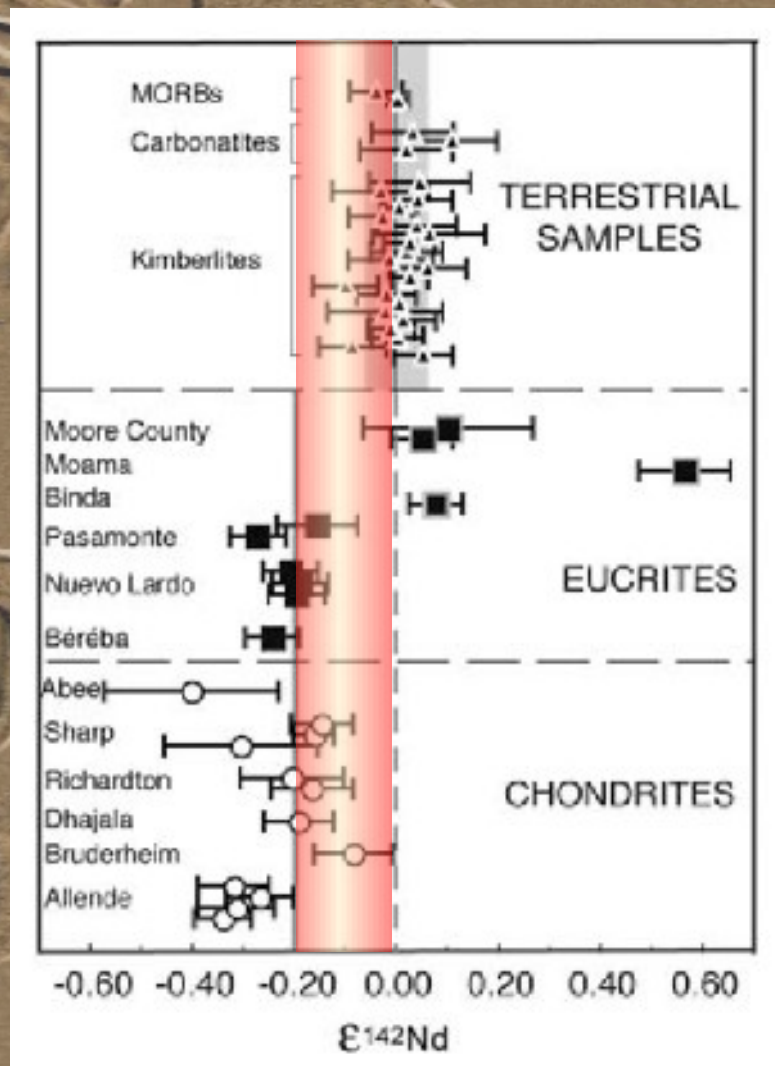
The End!



Différenciation précoce de la Terre à ~ 4.35 Ga en accord avec les anomalies en ^{142}Nd à Isua....

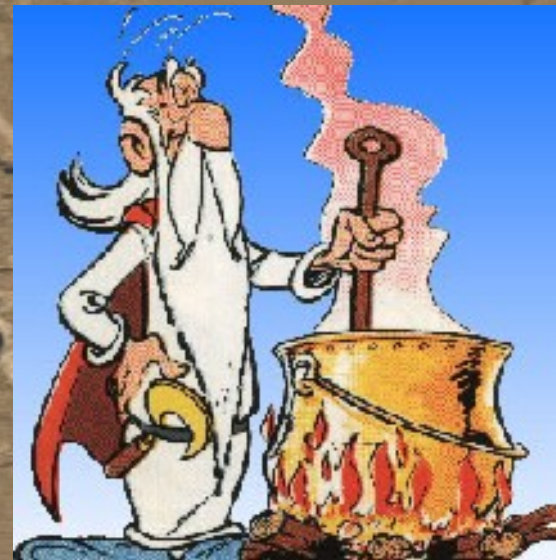


...et avec la différenciation majeure du manteau terrestre
~30 Ma après la fin de la nucléosynthèse

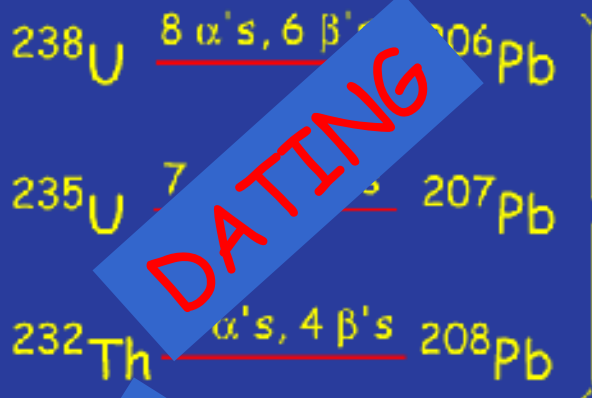




Méthodologie



ISOTOPIC SYSTEMS



DATING

TRACING



these have been widely exploited in SIMS

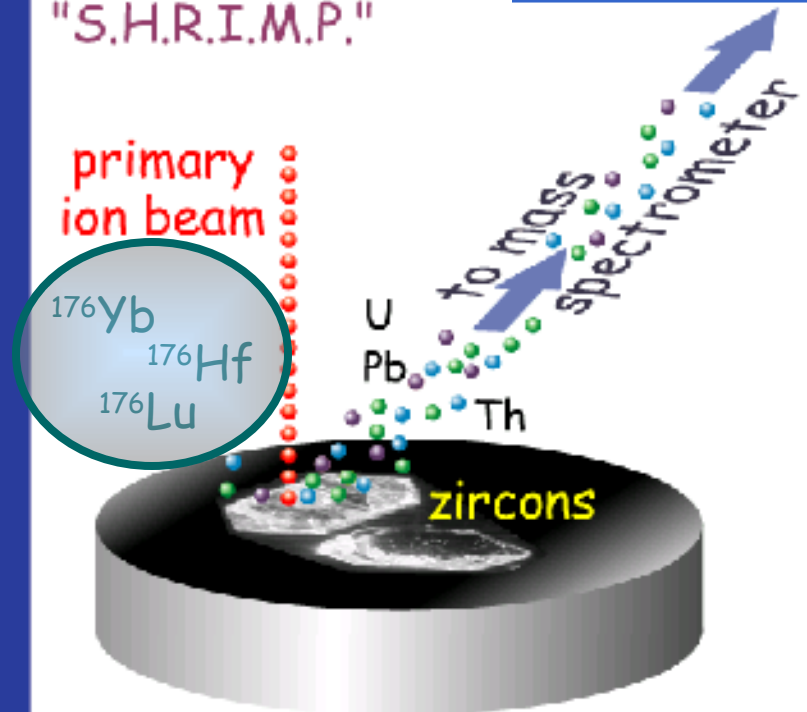


MC-ICP-MS

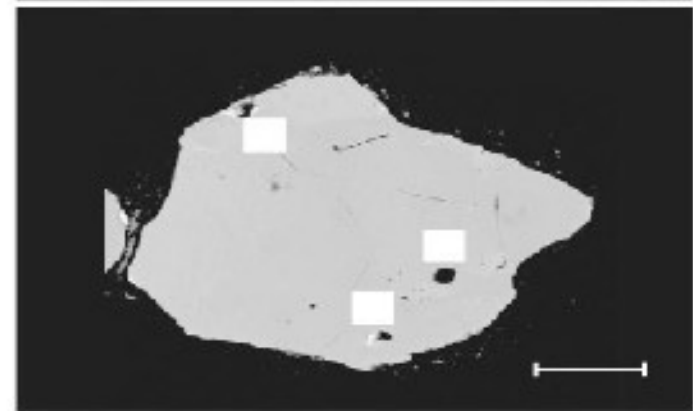
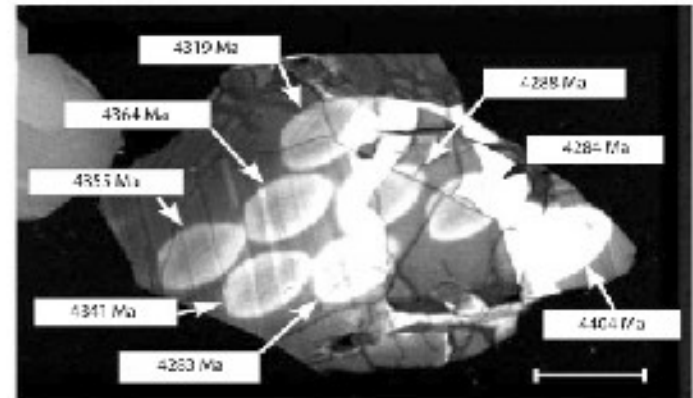
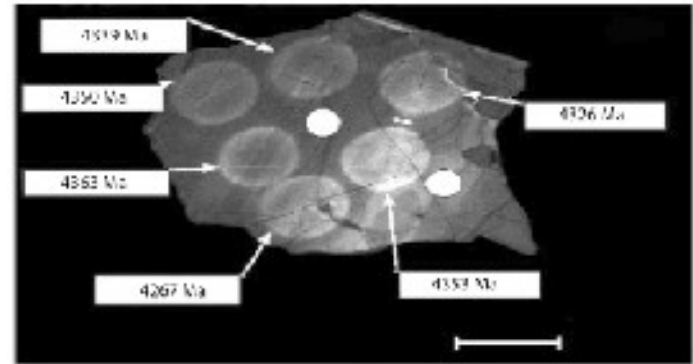
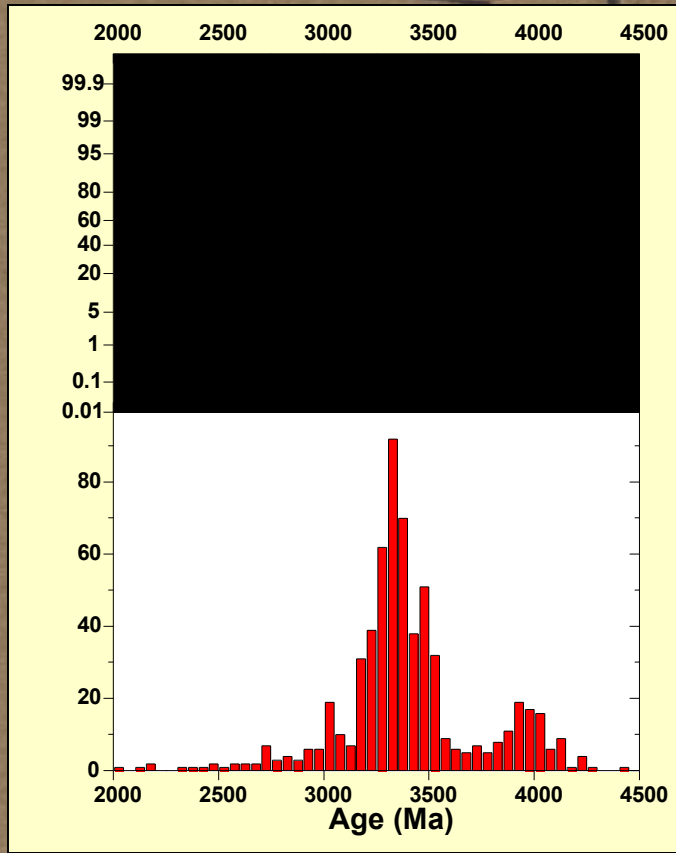
Beta (β^-) decay

SECONDARY ION MASS SPECTROMETRY

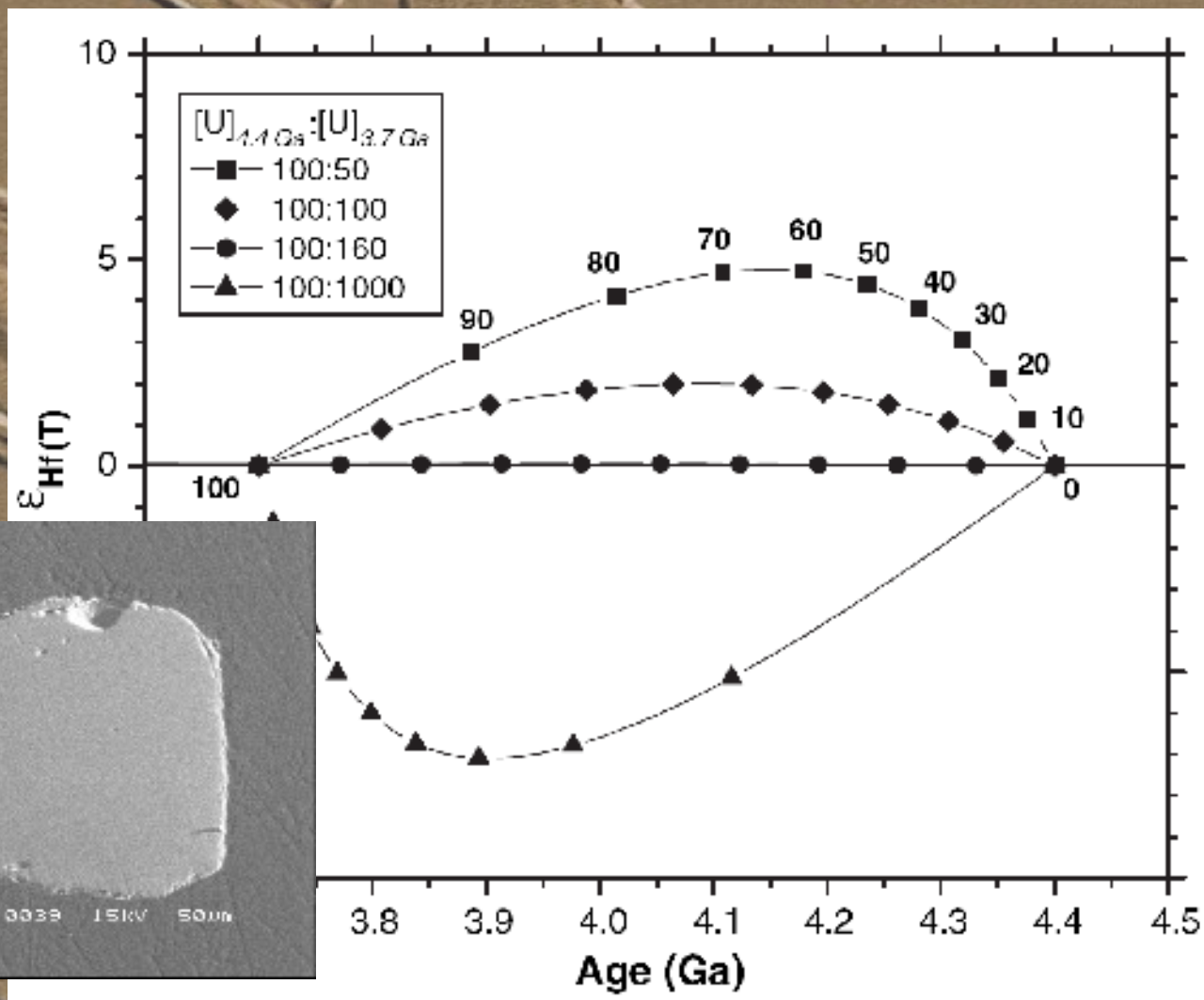
(SIMS or ion microprobe "S.H.R.I.M.P.")



The ion probe has made it possible to determine "in situ" ages on individual $\sim 20\text{-}\mu\text{m}$ spots on zircons and other minerals

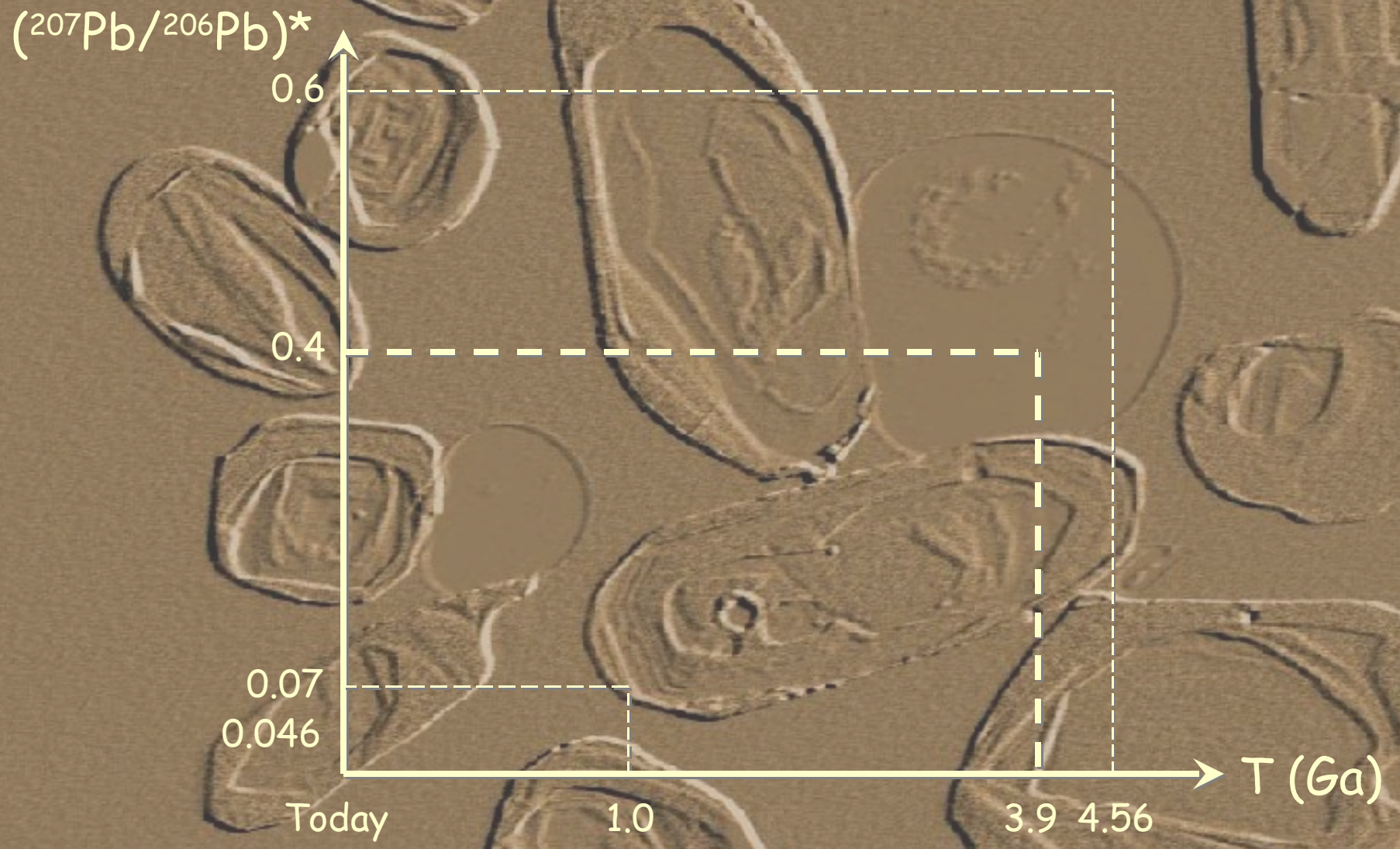


Mixing relationships among domains of different Hf isotope compositions and different U-Pb ages are non-linear



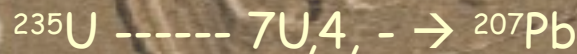
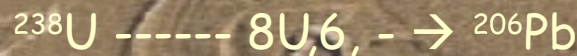
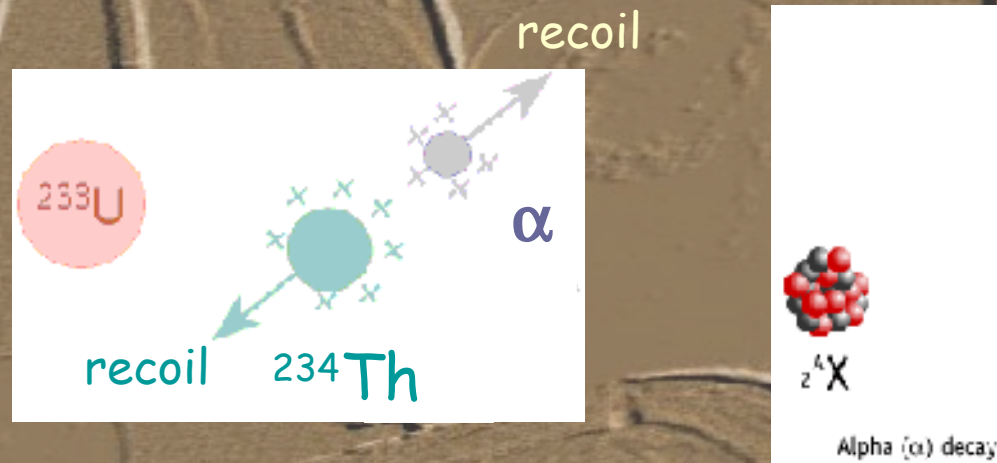
Pb-Pb dating

$(^{207}\text{Pb}/^{206}\text{Pb})^* > 0.4 \rightarrow \text{zircon} > 3.9 \text{ Ga}$

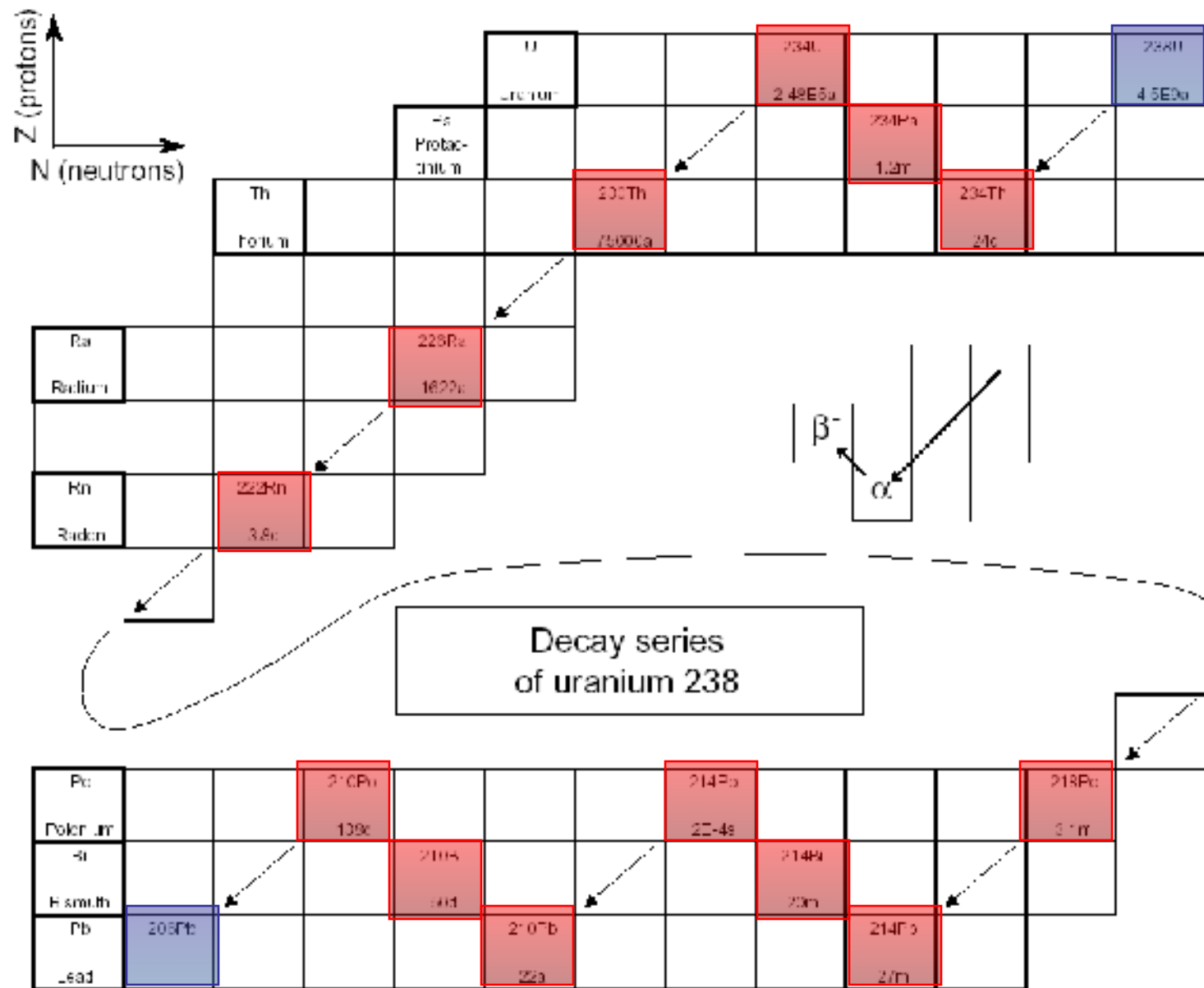


So what are the >4.25 Ga ages? Possibly recoil artifact ages!

....if crack present or zircon metamict (worse the higher the [U]) daughter lost...

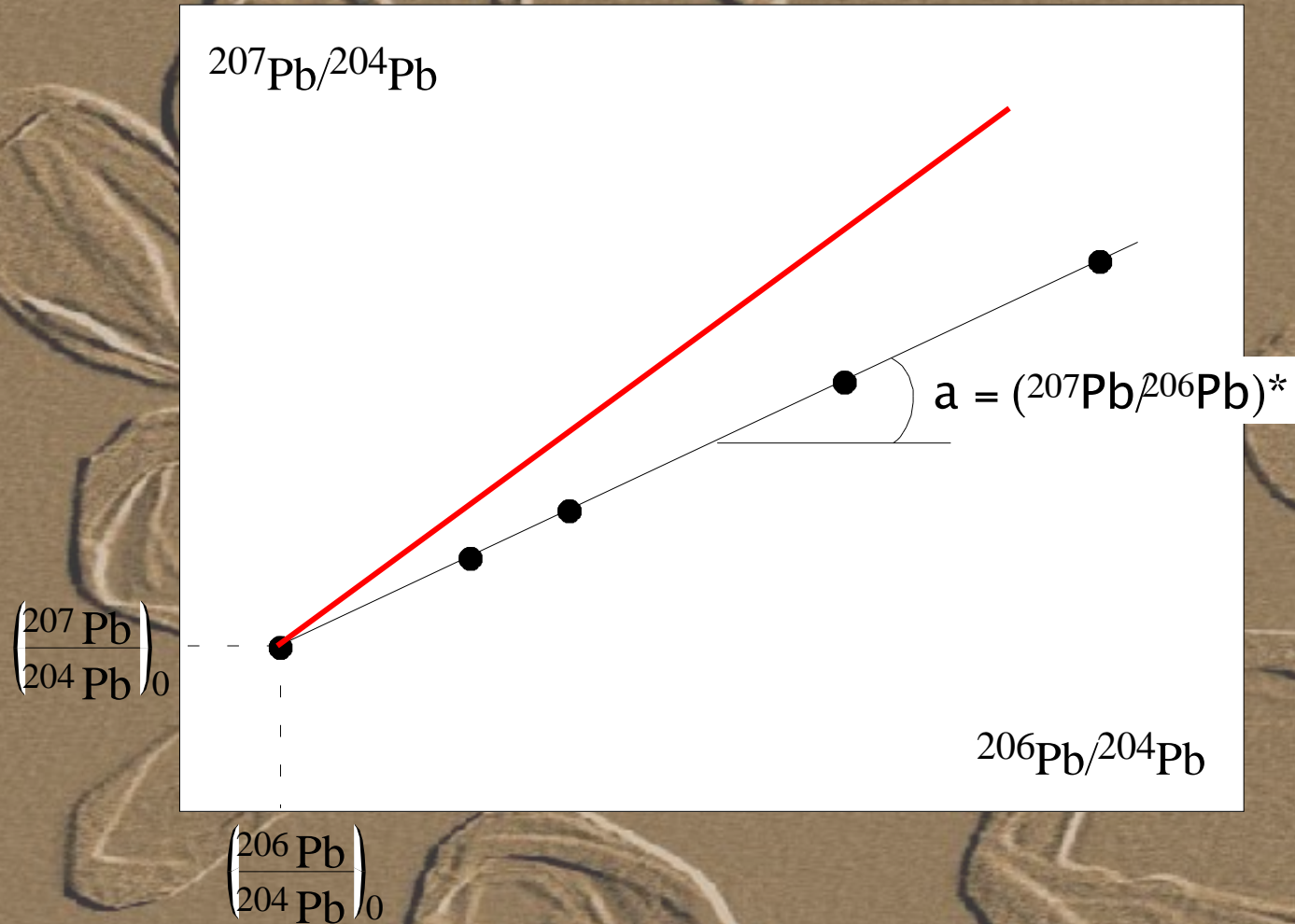


Recoil loss more extensive for $^{238}\text{U}/^{206}\text{Pb}$ system than for $^{235}\text{U}/^{207}\text{Pb}$ system due to fewer intermediate daughters for the latter → $^{207}\text{Pb}^*/^{206}\text{Pb}^*$ increases (or $^{206}\text{Pb}^*/^{207}\text{Pb}^*$ decreases)

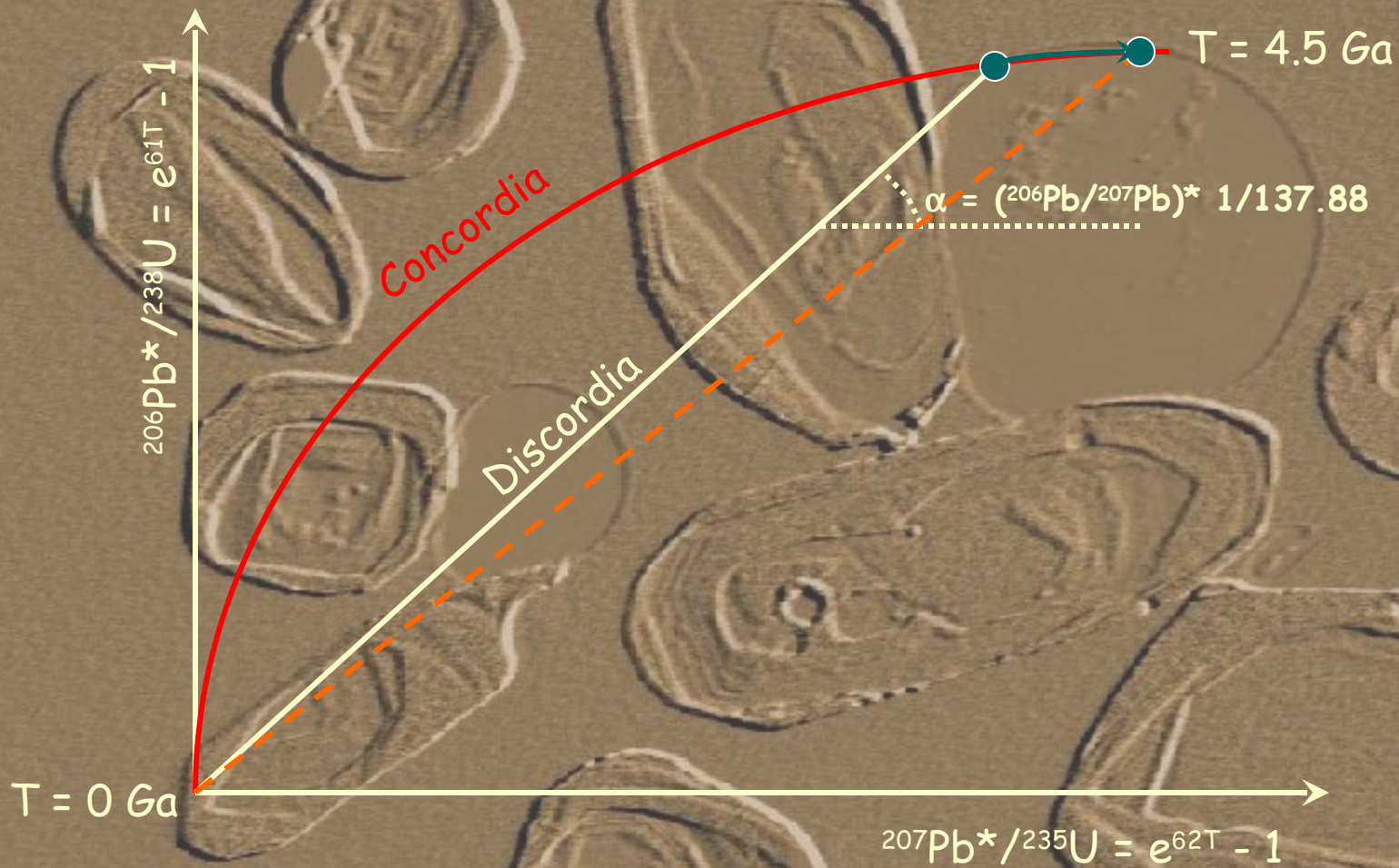


On Pb-Pb plot too high $(^{207}\text{Pb}/^{206}\text{Pb})^*$ gives
steep slope = too old age!

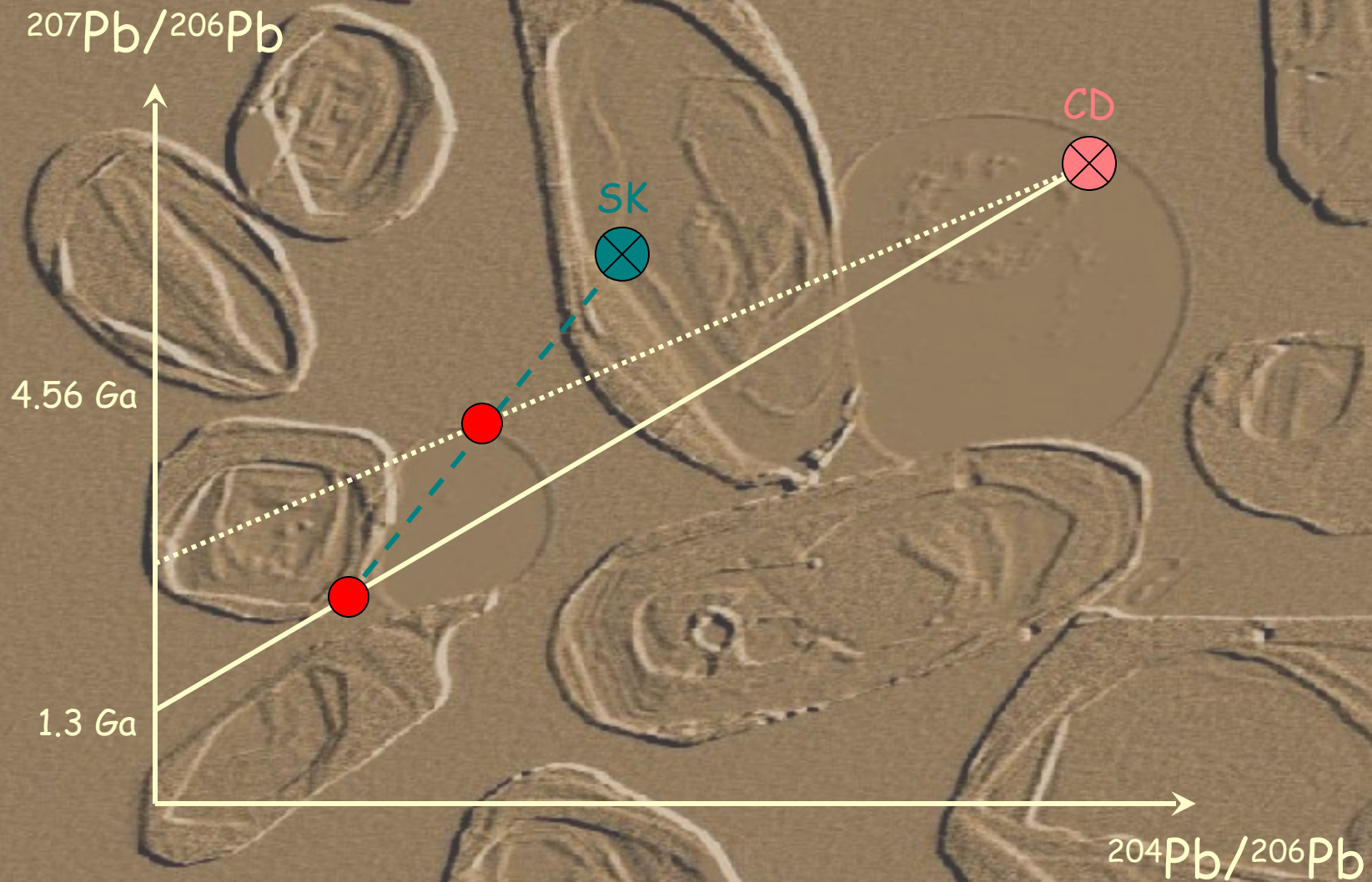
too

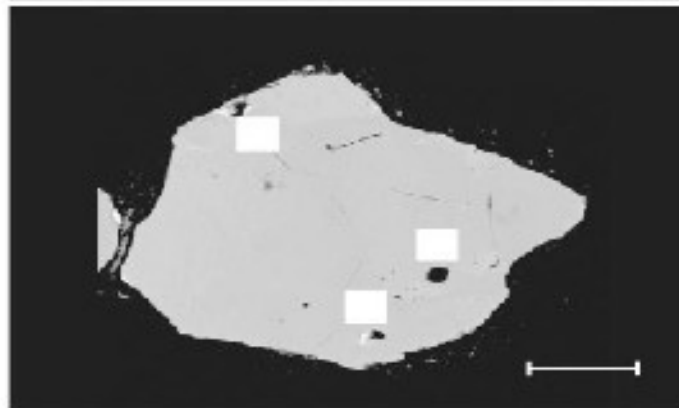
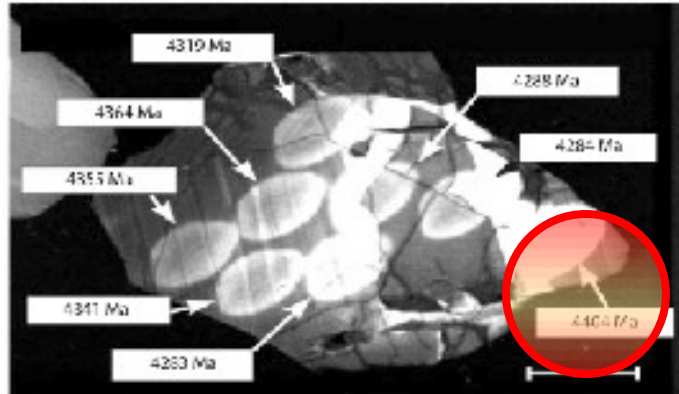
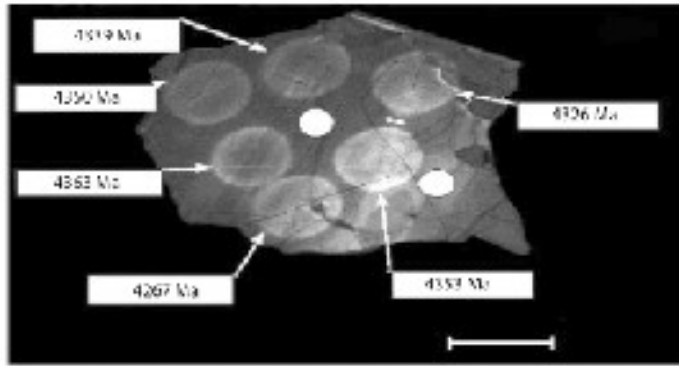


On Concordia plot too low $(^{206}\text{Pb}/^{207}\text{Pb})^*$ forces you up the Concordia towards too old an age!

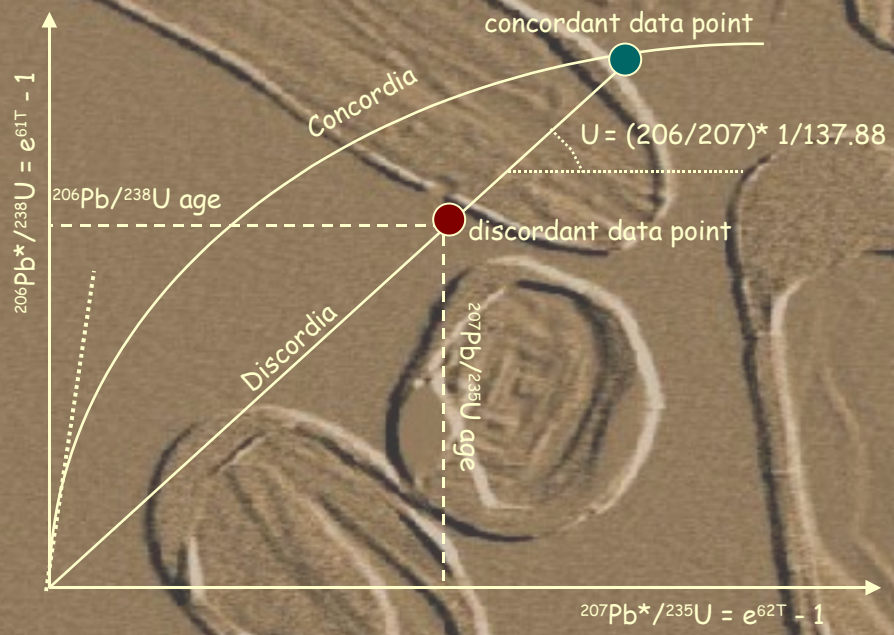


Or contamination by common Pb?





Appears upon
reinspection to be
right next to a crack!



Slope of Discordia = $\Delta y / \Delta x = \frac{({}^{206}\text{Pb}^* / {}^{238}\text{U}) / ({}^{207}\text{Pb}^* / {}^{235}\text{U})}{({}^{206}\text{Pb} / {}^{207}\text{Pb})^* \cdot {}^{235}\text{U} / {}^{238}\text{U}} = \frac{({}^{206}\text{Pb} / {}^{207}\text{Pb})^* \cdot 1 / 137.88}{({}^{206}\text{Pb} / {}^{207}\text{Pb})^*} = 1 / 137.88$

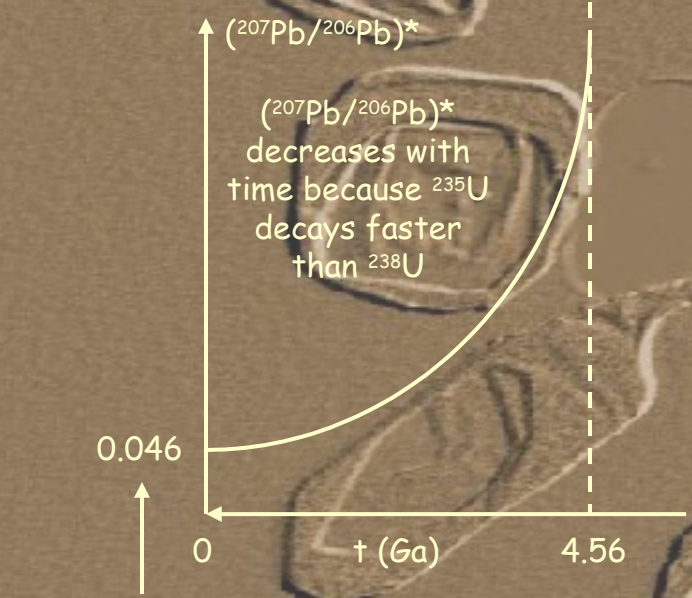
$({}^{206}\text{Pb} / {}^{207}\text{Pb})^* = 137.88 (\Delta y / \Delta x)$ or
 $({}^{207}\text{Pb} / {}^{206}\text{Pb})^* = 1 / [137.88 (\Delta y / \Delta x)]$

which compares with the slope of the Pb-Pb diagram in the following way:

$$a = \frac{1}{137.88} \left[\frac{e^{61t} - 1}{e^{62t} - 1} \right] = \left(\frac{{}^{207}\text{Pb}}{{}^{206}\text{Pb}} \right)^*$$



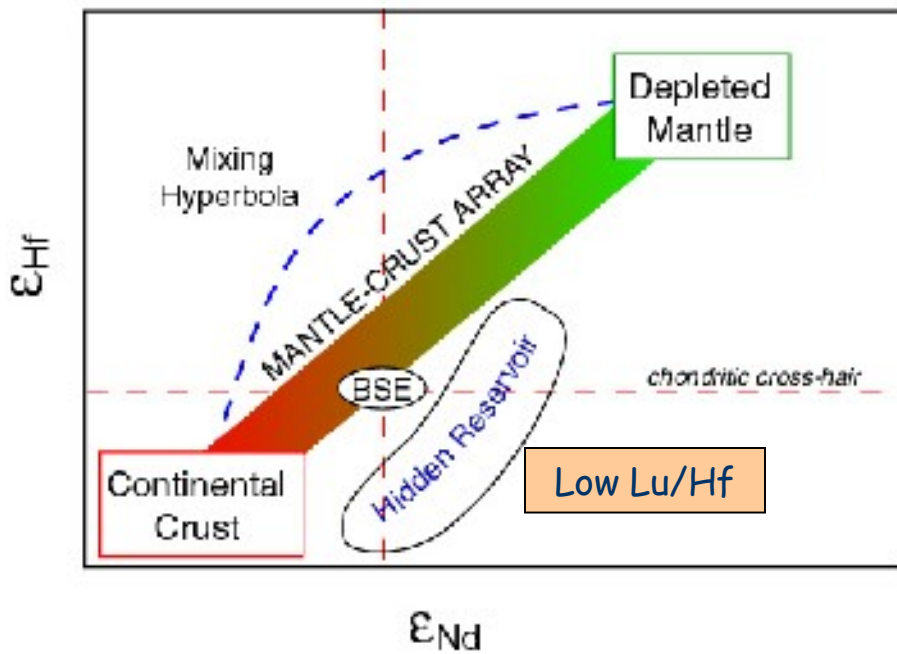
T Ga	y = e ^{61T} - 1	x = e ^{62T} - 1	λ1 (238U)	λ2 (235U)	(²⁰⁷ Pb/ ²⁰⁶ Pb)*
0	0.0000	0.0000	1.65125E-10	8.9485E-10	
0.2	0.00315	0.2177			0.05011
0.4	0.00540	0.4828			0.05470
0.6	0.0075	0.8055			0.05950
0.8	0.1321	1.1887			0.05580
1	0.1678	1.6774			0.07250
1.2	0.2046	2.2803			0.06012
1.4	0.2426	2.9701			0.05881
1.6	0.2817	3.8344			0.05872
1.8	0.3221	4.8869			0.11004
2	0.3630	6.1805			0.12289
2.2	0.4067	7.7202			0.13782
2.4	0.4511	9.6286			0.15183
2.6	0.4968	11.9437			0.17437
2.8	0.5440	14.7817			0.19682
3	0.5926	18.1637			0.22280
3.2	0.6426	22.1716			0.25242
3.4	0.6946	27.4597			0.28674
3.6	0.7480	33.6558			0.32835
3.8	0.8030	41.2004			0.37210
4	0.8598	50.3079			0.42501
4.2	0.9186	61.3753			0.48623
4.4	0.9789	75.1884			0.55714
4.6	1.0413	91.7873			0.63932



calculated from the slope of the Concordia at its origin

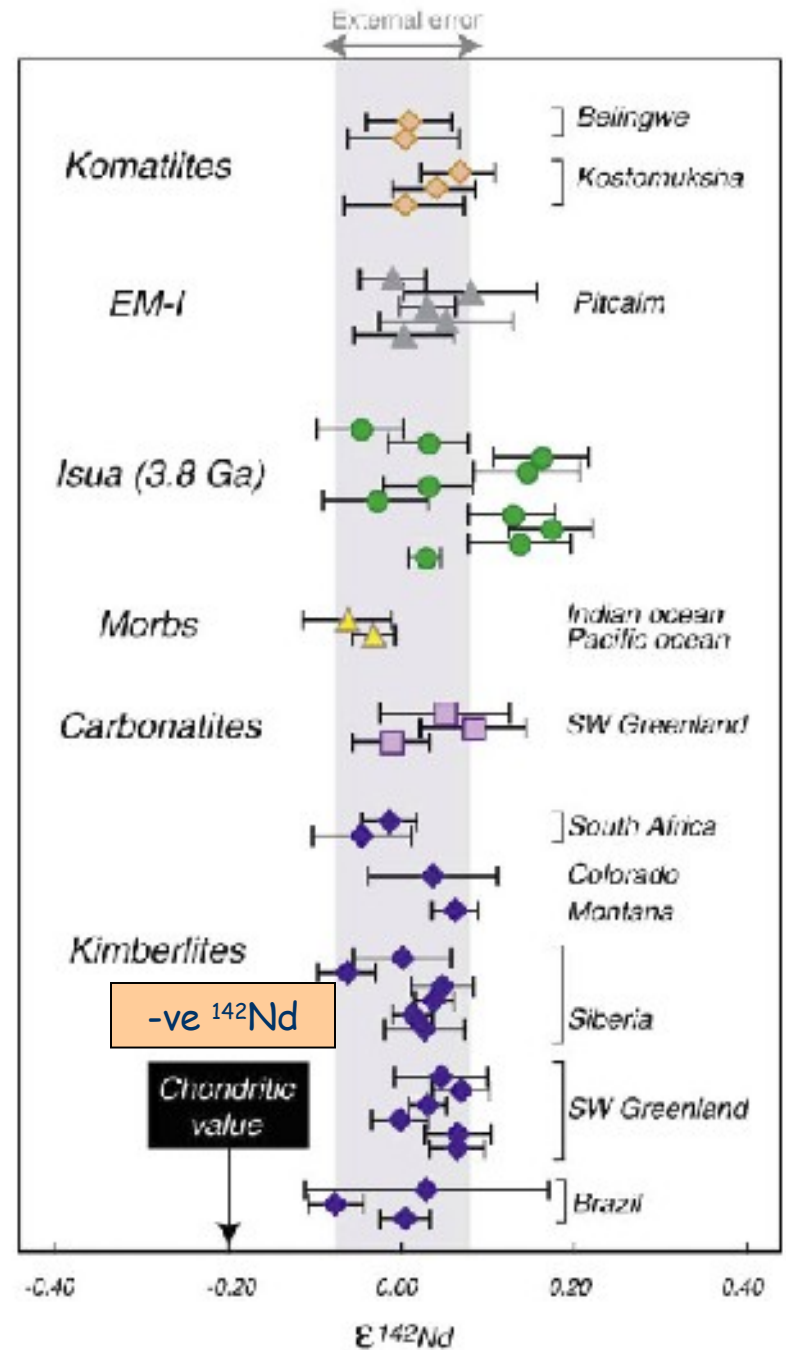
Reconnaissance SHRIMP II: (²⁰⁷Pb/²⁰⁶Pb)* ≥ 0.4 → older than 3.9 Ga

Evidence for a "hidden" reservoir from Hf and Nd isotopes in oceanic basalts and ^{142}Nd in chondrites

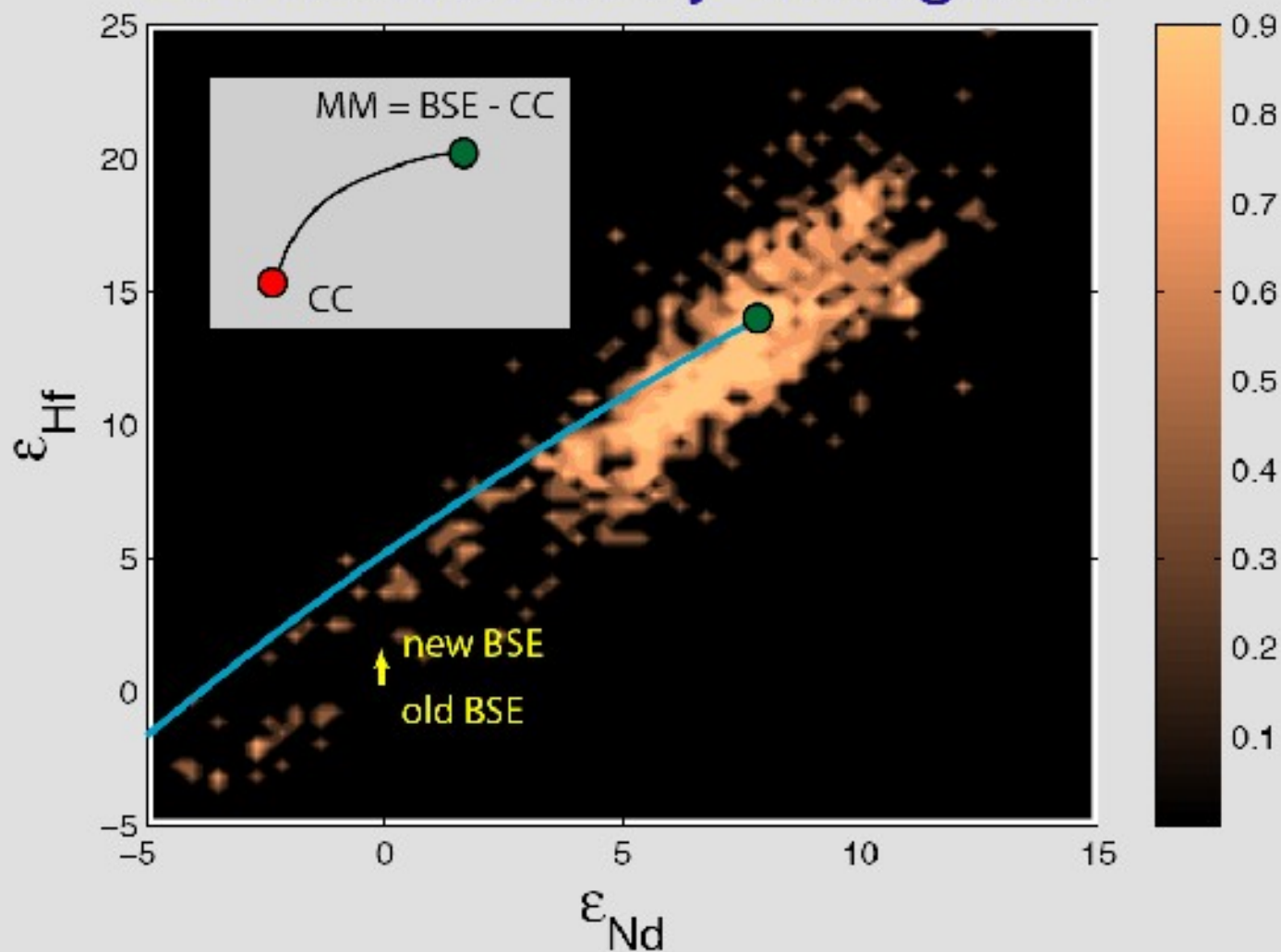


Blichert-Toft and Albarède (1997)

Boyet and Carlson (2006)



2D mantle array histogram



MM = mean mantle CC = continental crust BSE = bulk silicate Earth

